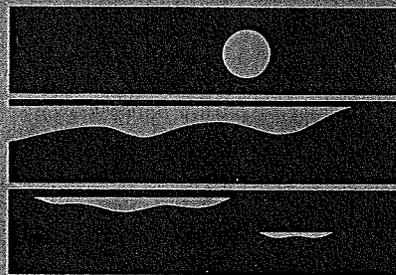


Water Resources Program

OPEN-FILE TECHNICAL REPORT



WASHINGTON STATE
DEPARTMENT OF
E C O L O G Y

DRILLING, HYDROLOGIC TESTING,
AND PIEZOMETER INSTALLATION
NEAR SINKING CREEK
LINCOLN COUNTY, WASHINGTON

D. B. Barnett

February, 1992

OFTR 92-1

This Open-File Technical Report presents the results of a hydrologic investigation by the Water Resources Program, Department of Ecology. It is intended as a working document and has received internal review. This report may be circulated to other Agencies and the Public, but is not a formal Ecology Publication.

Author: _____

D. B. Barnett

Reviewed by: _____

Linton L. Kildrick

Supervisor: _____

George E. Hanson

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ABSTRACT

During the summer of 1991, two NX-size (3-inch diameter) core holes and one shallow, 8-inch, auger hole were drilled at two locations within the Columbia River Basalts where Sinking Creek is adjoined by two springs (one active, the other inactive). The Wagner core hole penetrates the lower 141 ft. of the Priest Rapids Member basalt (Wanapum Formation), the Quincy interbed claystone (0.5 ft.), and the top 43 ft. of the Roza Member basalt (Wanapum Formation). The Rosman Core hole penetrates 10 ft. of Quaternary alluvium, 174 ft. of Priest Rapids basalt, 141 ft. of the Roza Member, 0.5 ft. of Vantage Member claystone interbed, and the top 9.5 ft. of the Grande Ronde Formation basalt. Core samples indicate that the Quincy interbed is absent at the Rosman site. The shallow auger hole penetrates 9.5 ft. of Quaternary alluvium, the lower 5 ft. of which is coarse gravel and sand.

Two short injection tests in the Rosman core hole and one pumping test of the nearby Baring Spring pool failed to demonstrate a direct hydraulic connection between the basalt aquifers and the spring. From the pumping test, a minimum transmissivity of 6000 ft²/day is calculated for the Quaternary alluvium.

One piezometer was installed in the Wagner core hole within the Priest Rapids basalt. Two piezometers were installed in the Rosman core hole; one each in the Priest Rapids Member and the Roza Member. One piezometer was screened in the gravel of the alluvium in the shallow auger hole. Water levels in the two Rosman core hole piezometers and the shallow piezometer indicated a downward, ground-water flow potential at this site, at the time of installation.

ACKNOWLEDGEMENTS

Many thanks to Art Larson of Environmental Investigations and Laboratory Services (Department of Ecology) and Linton Wildrick, of the Water Resources Program, for their careful review of the document.

INTRODUCTION

Ground water within the Columbia River Basalt Group (CRBG) of eastern Washington is the main source of water for crop irrigation in several counties, including Lincoln County, Washington. Since the late 1960s ground water levels in many of the agricultural areas of Lincoln County have experienced dramatic decline, especially during the irrigation season (Wildrick, 1982 and Olson, 1989). More recently, spring-fed, perennial and intermittent streams and lakes in these areas have shown a continuing depletion that is suspected to be a result of declining water tables or hydraulic heads in the basalt aquifers (Wildrick, 1982, and 1991). Sinking Creek, in northern Lincoln County, is one of these features.

This paper describes the methods and results of drilling, hydrologic testing, and piezometer installation in three test holes near two springs along Sinking Creek. Both springs are suspected to have hydraulic connection to the basalt aquifers, and do presently, or have in the past, contributed to flow in Sinking Creek.

Purpose of the study

Three main purposes of the project were: 1. To test for hydraulic interconnection between springs and basalt aquifers through hydrologic testing conducted in drill holes; 2. To monitor water levels in the basalt aquifers during successive cycles of irrigation by installing piezometers in the holes. This will allow long-term analysis of pumping effects on the surface and subsurface hydrologic regimen, and; 3. To gather reliable stratigraphic and lithologic information from the drilling for local and regional hydrostratigraphic correlation.

Location of the study area

The study area is in northern Lincoln County, Washington, near the northern edge of the Columbia Plateau (or *Columbia Basin*, as described by Freeman and others (1945)) physiographic province, and just south of the towns of Wilbur and Creston (Figure 1).

The Wagner core hole (Figure 2, Plates 1 and 2) is on property owned by Clarence Wagner and son Dan, who use the land for ranging cattle and growing wheat and hay. The drill hole is approximately 100 ft. northeast of the extinct "Wagner Spring" (T25N/R35E-12K1s). Land surface elevation of the Wagner hole is approximately 2090 ft. above mean sea

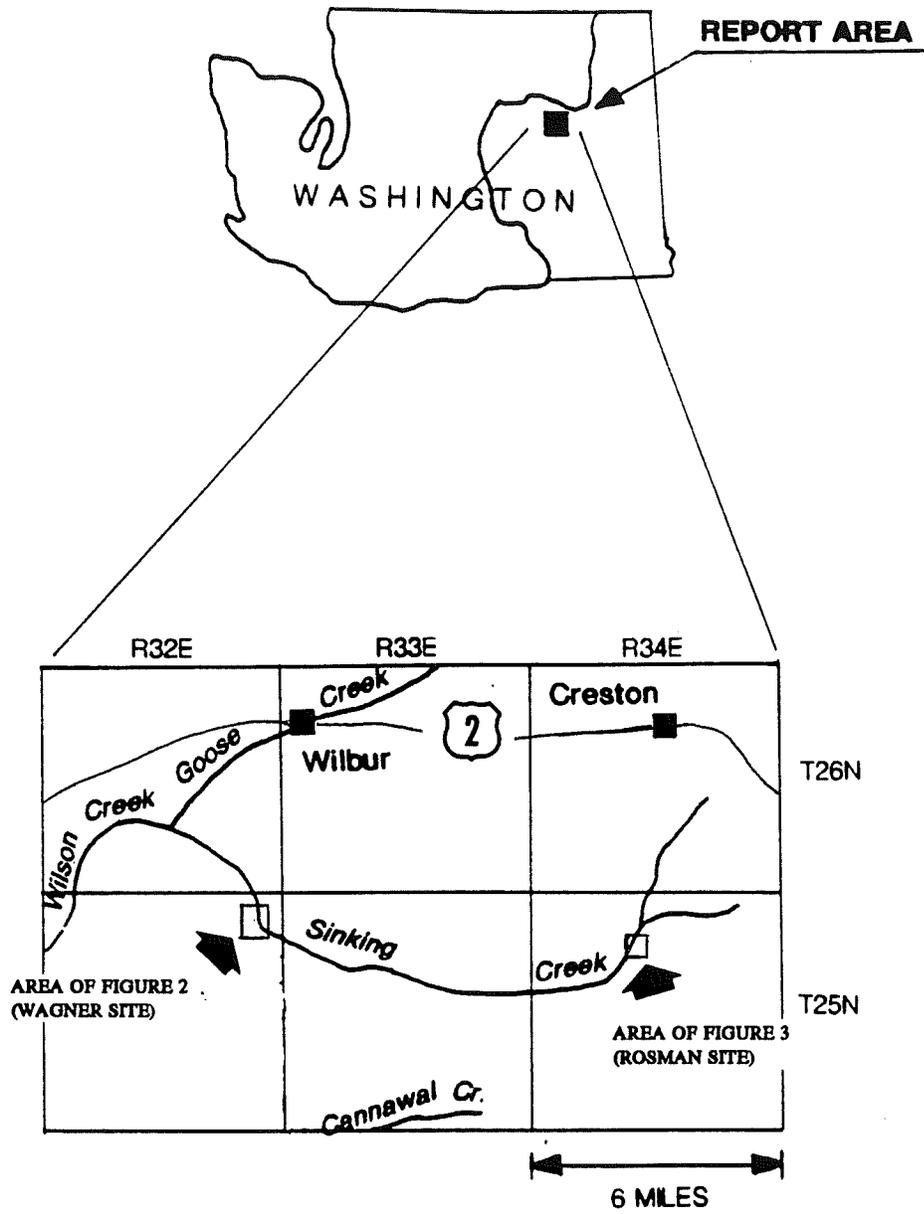


Figure 1. Location of drill holes



Plate 1. The Wagner site, looking south. The rock-lined, dry channel of Sinking Creek appears in the middle distance. Arrow denotes location of the extinct Wagner Spring.



Plate 2. View looking northwest along the channel of Sinking Creek from the Wagner drilling site. Fragments of a dilapidated concrete containment structure, built for the spring, are visible in the foreground.

level (estimated from U.S. Geological Survey, Wagner Lake 7.5 minute quadrangle).

The Rosman core hole (Figure 3 and Plate 3) is 50 ft. northeast of Baring Spring (or "Rosman Spring" as called by local residents) (T25N/R34E-8N1s) on land owned by John Rosman. The surrounding land is range and hay fields. The site occupies the northern edge of the flood plain of Sinking Creek, roughly 50 ft. from an outcrop of Priest Rapids basalt that marks the northern side of a shallow coulee. At this site, the coulee is roughly 500 ft. wide. Land-surface elevation of the Rosman site is approximately 2210 ft. above mean sea level (estimated from U.S. Geological Survey, Creston Butte 7.5 minute quadrangle). The U.S Geological Survey map erroneously shows Baring Spring 3000 ft. west-southwest of its actual location.

Previous work

The most complete compilations of regional geology, including the Sinking Creek area, are those of Waggoner (1990c) and Stoffel and others (1991). Swanson and others (1979a) focused geologic mapping efforts on identifying basalt flows. Reidel and Hooper (editors, 1989) present the most recent compendium on stratigraphic, structural, and petrologic aspects of the CRBG.

Gephart and others (1979) discuss the factors controlling movement of ground water in the CRBG. Wildrick (1982, 1985, 1990, and 1991) assesses earlier hydrologic studies, well data, aquifer tests, and geology in the Sinking Creek drainage to evaluate the interconnection between basalt aquifers and springs.

Pachis and others (1988), Steele and others (1988), and Tera Corporation (1983) report on the results of drilling and hydrologic testing in basalts, for the purposes of characterizing contaminant transport in fractured basalt and assessing general hydrogeologic conditions in the CRBG. The sites of these investigations are from slightly one mile south-southeast of the Rosman site, to a few miles northeast thereof.

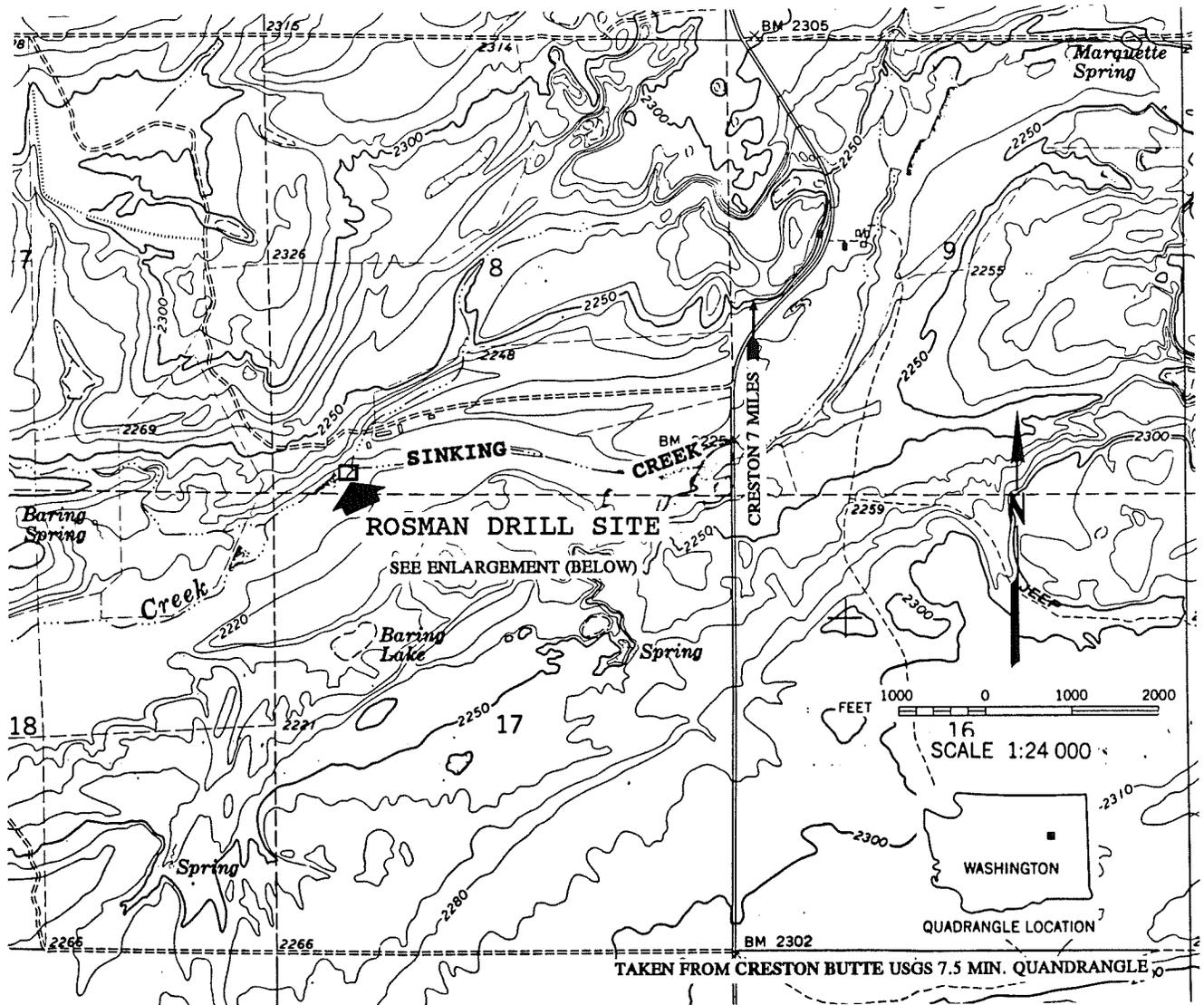


Figure 3.
Rosman site

DETAIL OF AREA ENCLOSED ABOVE

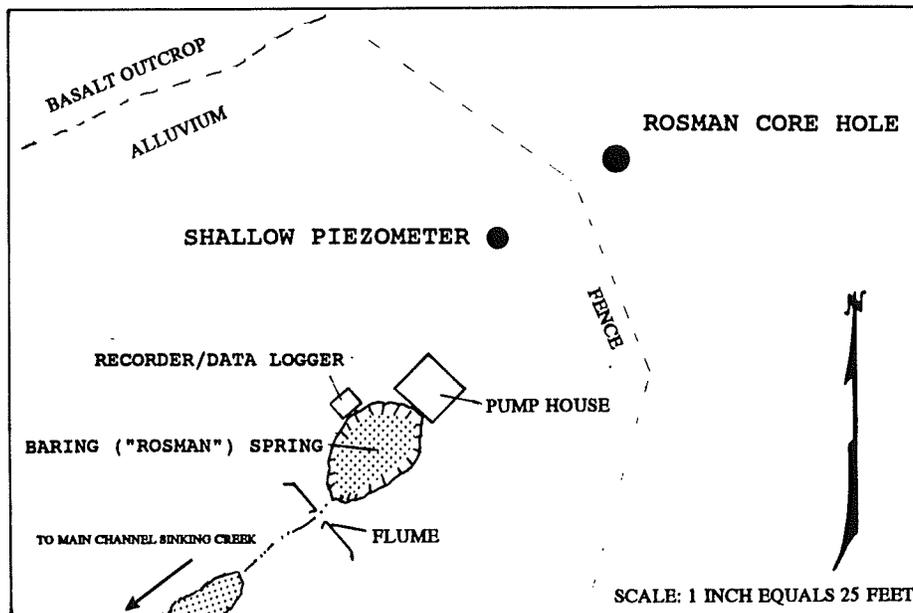




Plate 3. The Rosman site. The photo was taken from the basalt outcrop, looking southeast across the coulee. The pump house for the stock-watering pump and the instrument shelter (left center) are on the edge of the spring pool. An artificially-straightened portion of Sinking Creek appears in the middle distance (outlined, in part, by the line of bushes). The Rosman shallow piezometer hole was drilled near the rear tire of the backhoe. The Rosman core hole was drilled just left of the backhoe on the far side of the wooden fence. The flume shown in Figure 3 had not been installed at the time of the photograph. Note the light-colored pile of diatomite-rich alluvium next to the instrument shelter.

STRATIGRAPHY AND HYDROGEOLOGY

The project area lies near the northern boundary of the lava flows of the Miocene Columbia River Basalt Group (CRBG). Wildrick (1991) summarizes the relationship between stratigraphy and aquifer occurrence in the Sinking Creek area, and describes the nature of geologic materials supporting the creek channel. Figure 4 illustrates the general stratigraphy of the project area. The schematic, as-built diagrams of Appendix A shows the general stratigraphy at the drill sites.

The oldest rock encountered in the drill holes is the Grande Ronde Formation basalt, generally described as a dense, black, mostly aphyric basalt. This unit outcrops several miles to the north of the project area, and was encountered in the bottom 9.5 ft. of the Rosman core hole.

The Vantage Member sedimentary interbed of the Ellensburg Formation separates the Grande Ronde Formation from the overlying Wanapum Formation. In the Rosman hole, this unit occurs at 325.0 ft. below land surface (bls) and is a 0.5-foot-thick, soft, dark gray-green claystone. The Wagner hole is not deep enough to intercept the Vantage Member.

The Roza Member basalt of the Wanapum Formation directly overlies the Vantage Member interbed, and is exposed in the deeper canyons near the Columbia River, north of the project area. The Roza basalt is mostly aphyric in groundmass, but with striking plagioclase phenocrysts up to 1 cm in length that serve to easily identify the member. In the Rosman core hole, some intervals display a glomerophyric texture. The Roza Member was encountered at 142.4 ft. bls to total depth (185.4 ft. bls) in the Wagner core hole. The unit is 141 ft. thick (184.0 ft. to 325.0 ft. bls) at the Rosman site.

Directly above the Roza Member is the Quincy sedimentary interbed. In the project area, the Quincy interbed is usually a thin (1-2 ft.) claystone, is not ubiquitous, and may vary considerably in thickness over short distances. The Quincy is a 0.5-foot-thick, soft, greenish-brown claystone in the Wagner core hole, but is absent in the Rosman core hole.

The Priest Rapids Member of the Wanapum Formation overlies the Quincy interbed horizon, and is the only basalt unit exposed in the study area. The Priest Rapids Member is a medium-dark gray, microphyric basalt. A dense field of interwoven plagioclase laths is barely discernable in hand specimen. Outcrops of the Priest Rapids Member occur near both drill sites.

		Valley Alluvium	Catastrophic Flood Gravels	
Columbia River Basalt Group	Wanapum Basalt	Priest Rapids Member		Unconfined
				Aquifers
		Quincy Interbed		Confining Bed
		Roza Member		Upper Confined Aquifer
	Grande Ronde Basalt	Vantage Interbed		Confining Bed
Member Uncertain		Lower Confined Aquifer		

Figure 4. Hydrostratigraphy for the Sinking Creek Area (Modified after Wood, 1987b).

Loess (mostly eolian silt), deposited from late Pliocene through Holocene, overlies the CRBG and is intercalated with catastrophic flood deposits. The largest accumulation of this material is known as the Palouse Formation of "Bull Lake" age (200 ka to 125 ka) (Baker and Nummedal, 1978).

During and following latest stades of Pleistocene glaciation, large channels (or *coulees*) were incised into the basalt and surficial sediments by the erosive force of multiple, catastrophic floods, collectively known as the "Spokane Flood" (Baker and Nummedal, 1978; McDonald and Busacca, 1988). Gravel and sand deposits, left from aggradational stages of the flood, partially fill the coulees in places and mantle some of the higher terrain. Late Pleistocene or Holocene (post-flood) lake deposits cover much of the coulee bottoms. Lake deposits consist of clay, silt, volcanic ash, peat, and diatomite. Elsewhere, post-Flood loess and remnants of pre-flood loess overlie and are interbedded with both flood and lake deposits.

Hydraulic characteristics of basalt

Ground water moves principally through subvertical and subhorizontal fracture systems within basalt flows and within subhorizontal zones of higher permeability at the bases and tops of the flows (Gephart and others, 1979). Fractures originate from cooling of the basalt lava or from later, tectonic deformation. The hydraulic conductivity of fracture systems is related to the intensity of fracturing, the connectivity of the fractures, and the degree of blockage by secondary, in-filling materials, such as clay.

Fault zones in basalt may simultaneously transmit and confine ground water (Gephart and others, 1979). The plane of the fault and slip surfaces will often host gouge and impermeable clay, but the associated, peripheral fracture system may be highly permeable. Thus, fault zones and fracture systems may be sites of considerable vertical leakage of ground water and/or boundaries to horizontal flow (Newcomb, 1961 and 1969, and Pachis and others, 1988).

METHODS OF STUDY AND RESULTS

BASIS FOR SITING DRILL HOLES

Both core holes were located within a few feet of active or extinct springs and the main channel of Sinking Creek. Through the study of the stratigraphy and observation of ground-water levels near the spring sites, we hoped to gain insight into the mechanisms of interconnection between springs and basalt aquifers in the area.

The Wagner hole was drilled within a few feet of the defunct "Wagner Spring", on the banks of Sinking Creek (Plates 1 and 2). The spring had contributed to flow in the creek before 1977 (Clarence Wagner, personal communication). The main objectives at this site were to gather detailed lithologic information from the drilling, and to install at least one piezometer as a local ground-water monitoring point within the Priest Rapids basalt.

At the Rosman site (Figure 3, Plate 3), Baring Spring was not flowing during the time of the project, but maintained a pool that fell steadily during the summer. The spring normally flows during winter and spring. If, as postulated, the basalt aquifers feed this spring, a sufficient hydraulic stress, imparted within a core hole in the basalt several feet from the spring, might induce a measurable response in the pool of the spring. This relationship might be especially notable if the interconnection between spring and aquifer(s) were via a two-dimensional, highly-conductive "conduit" as suspected by Wildrick (1985). A similar structure was detected in hydrologic tests in a nearby well field by Steele and others (1988). Testing would consist of isolating specific intervals of the core hole, in the basalt, with an inflatable packer and injecting a metered amount of water into the isolated interval while monitoring the level of the spring pool.

Alternatively, if no distinct structure exists, leakage between aquifers may be occurring over a large area. If significant, upward or downward flow potential exists in the proximity of the spring, properly-placed piezometers would define this. Wildrick (1991) more recently favors the possibility that basalt aquifers may contribute to, or receive ground water from the broad alluvial aquifer that hosts the spring. Pumping of the spring pool might induce a response in the nearby piezometers.

A second, shallow hole was drilled to near the base of the alluvium at the Rosman site to observe the behavior of the water-table aquifer in the alluvium, both during the

injection testing in the Rosman core hole, and during the pumping test of the spring.

DRILLING

All three holes were drilled with a track-mounted, high-speed wireline diamond drill using a "Hydracore 28", Longyear Series 210, hydraulic operating system. Both the Wagner and Rosman core holes were drilled with "N"-size rods and "NQ" diamond bits. Nominal diameters of the holes and core are 2.98 inches and 1.875 inches, respectively.

Surficial materials (alluvium) were drilled to refusal with a hollow-stem auger bit. Drive samples of fine-grained sediments were taken through the auger bit to a depth of 6 ft. bls at the Rosman core hole.

All core was logged in detail and boxed for storage. Particular attention was focused on the number, orientation (with respect to horizontal), width, and filling materials of fractures. Original, lithologic logs for each hole appear in Appendix B.

Steel surface casing was emplaced using the rotary action of the drill and a diamond-impregnated casing shoe. Surface casing for the "N"-size hole has the same nominal dimensions as "H"-size drilling rods (3.5 inches, O.D.), and is essentially reamed into place within the existing "N"-size hole. Project leaders initially intended to remove the surface casing upon completion of the holes, but problems with the drilling equipment made it necessary to cement the casing in place in the Rosman core hole.

A summary of the progress and highlights of each hole follows.

Wagner core hole

The Wagner core hole was spudded on June 11, 1991 using 4-inch and 8-inch auger bits. The hole was drilled in this manner to a depth of 8 ft. bls. Wireline coring methods were then used to drill to a total depth of 185.4 ft. bls. Total depth was reached June 19, 1991. Temporary, steel surface casing, with a diamond shoe, was installed to a depth of 12.0 ft. bls.

The average drilling rate for the Wagner hole was 20.6 ft./day, but days of greater than 30 ft. were common when the rock was competent. Coring runs were typically 4-5 ft. in length (inner tube capacity = 5 ft.), but occasionally

less than one foot in fractured basalt. Severely fractured basalt invariably caused jamming of the core barrel and slow drilling. Significant fracturing and resulting slow drilling was encountered in the upper 50 feet of the hole, and in 3 other intervals deeper in the hole, each roughly 5 ft. in length.

Eight feet of loess and alluvium were penetrated at the Wagner site before the Priest Rapids Member basalt was reached. The Priest Rapids Member is a medium-to-dark gray basalt, with widely dispersed vesicles and amygdules 1-2 mm in diameter that diminish in abundance downward. The vesicles and amygdules disappear completely at 77.5 ft. bls, and the rock becomes massive and relatively unfractured. Large vesicles, often in discrete bands or layers, reappear at 123.0 ft. bls, and accompany an increase in fracture abundance. Scoriaceous basalt (flow bottom?) and flow breccia are encountered at 130.6 ft. to 133.9 ft. bls and at thinner horizons down to 141.7 ft. bls.

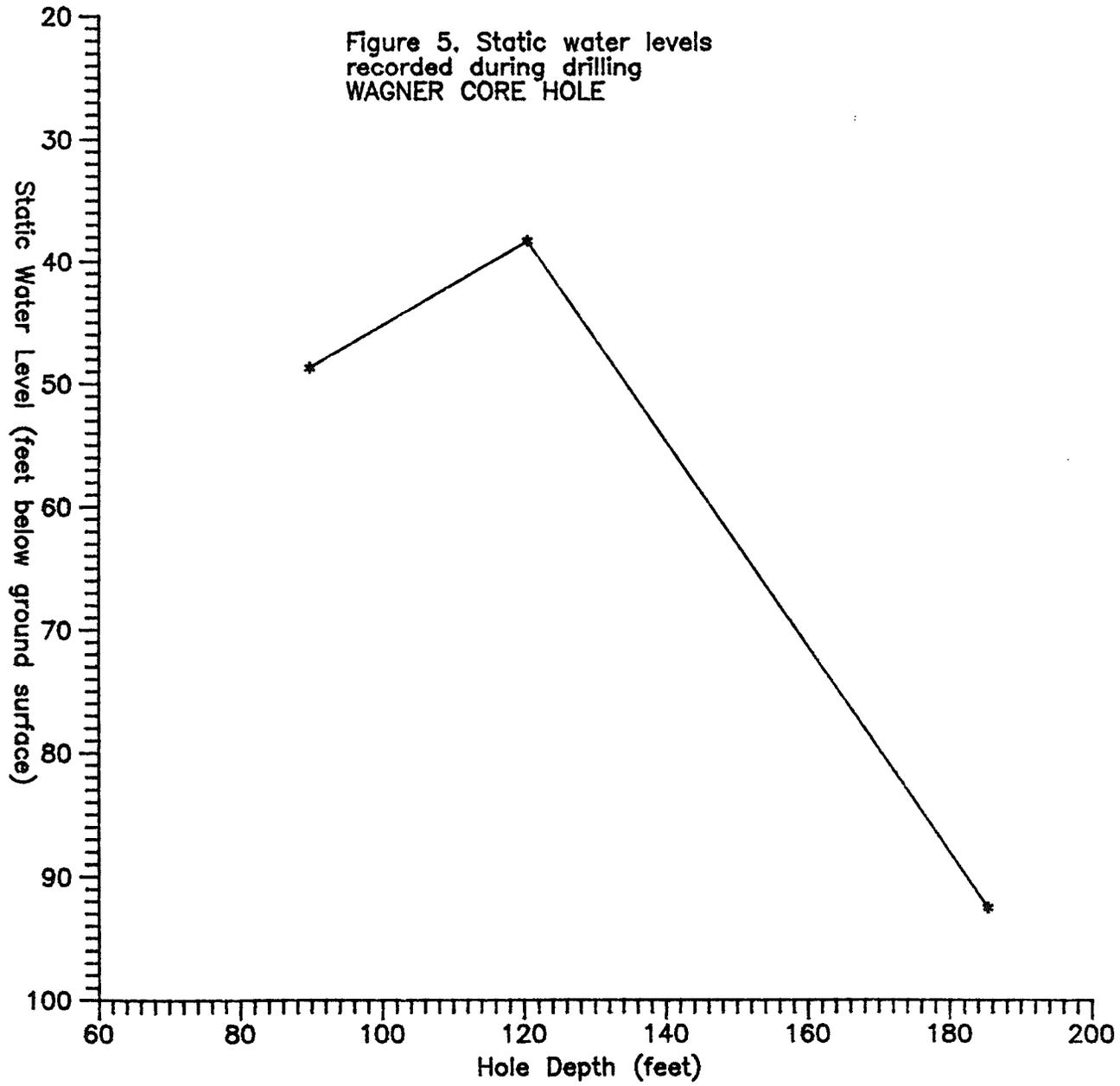
The Quincy Interbed, is found at 141.7 ft. to 142.4 ft. bls. At this location, the Quincy is a soft, moderately-friable, greenish-brown claystone.

Vesicular, porphyritic basalt of the Roza Member lies directly below the Quincy Interbed. Vesicles in this unit also tend to occur in distinct bands after passing through a uniformly-vesicular flow top. Drilling ceased after penetrating 43.0 ft. of the Roza Member.

Fracture-filling or fracture-coating materials in the basalt of the Wagner hole are mostly iron oxides (hematite, goethite(?), and limonite) and clay. Other minerals identified in hand specimen were pyrite, zeolite(s), manganese oxide, calcite, and possibly chlorite.

Ground water was first encountered between 42.7 ft. and 89.8 ft. bls in the Wagner hole. Because of large amounts of water used during drilling, static-water-level measurements were made only in the morning, prior to drilling activities for the day. On the morning following the last day of drilling, the static water level was 79.98 ft. bls. Figure 5 illustrates static water-levels recorded during the drilling of the Wagner hole.

Figure 5. Static water levels recorded during drilling WAGNER CORE HOLE



Rosman Core hole

Drilling began for the Rosman core hole on June 21, 1991. An 8-inch hollow-stem auger was used to drill through the first 10.2 ft. of alluvium. Continuous drive samples were taken from the surface to 6.5 ft. bls, where gravel prohibited any further progress with the sampling tube. Wireline, NX-size core drilling was used from 10.2 ft. to the total depth of 335.2 ft. bls. NX-size surface casing and a diamond casing shoe were installed to a depth of 33.5 ft. bls into Priest Rapids Member basalt. Total depth was reached on July 15, 1991.

Drilling rates averaged 25.8 ft./day, discounting days of testing or of no drilling activity. Many days saw over 40 ft. of drilling. Drilling rates were enhanced when the drillers acquired a new, 10-foot inner tube for holding core, so fewer wireline trips were necessary. Although 10-ft. runs were possible with the new equipment, intervals of fractured basalt caused core-barrel jamming and short runs at many levels (Plate 4). Significant intervals of fracturing occur from; surface to 30 ft. bls, 150 ft. to 177 ft. bls, and 280 ft. to 325 ft. bls.

From the surface, five feet of clay-rich soil and diatomite overlies 5-7 ft. of gravelly alluvium. The lower few feet of the gravel is extremely coarse, and highly conductive of ground water. The base of the gravel lies directly upon highly-fractured basalt of the Priest Rapids Member at 12.4 ft. bls.

The Priest Rapids basalt is almost uniformly vesicular/amygdaloidal throughout, with the exception of isolated horizons of higher vesicle density at widely-separated intervals. Vesicles almost completely disappear from 139.0 ft to 180 ft. bls. The vesicles reappear abruptly at 180 ft., and continue down to the top of the Roza Member basalt (184.0 ft. bls).

The Quincy interbed is absent at this site. In its place, is a highly-vesicular rubble zone representing the flow-top material of the Roza Member basalt.

The Roza Member basalt is encountered from 184.0 ft. to 325.0 ft. bls. The rock is distinguished by its characteristically large plagioclase phenocrysts. Zones of significant fractures in the Roza Member occur from: 225.0 to 236.4 ft. bls, 259.0 ft. to 264.3 ft. bls, and 278.0 ft. to 325.0 ft. bls (bottom of Roza). Horizons

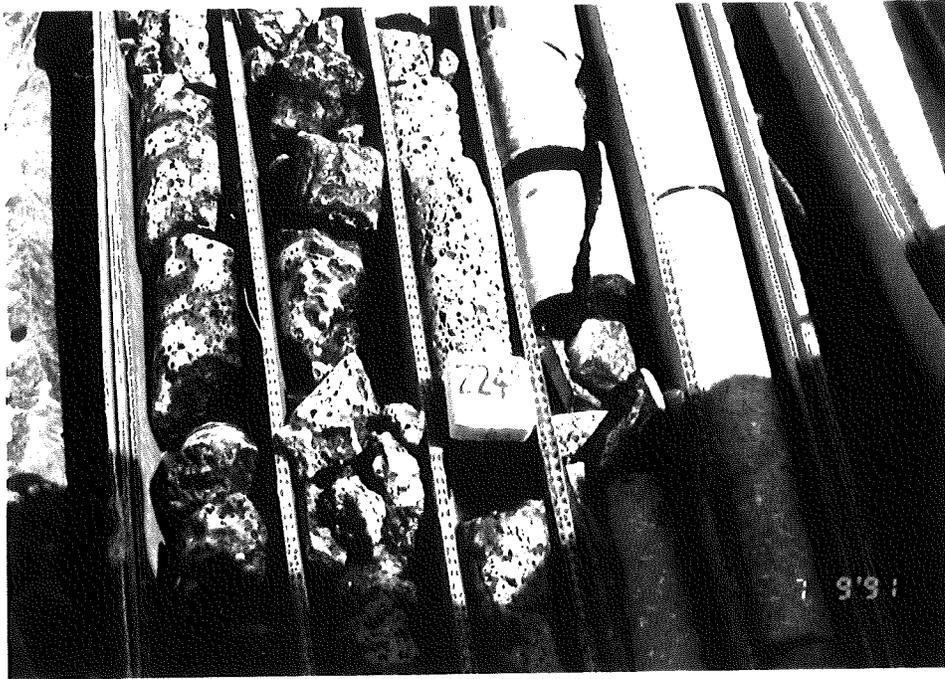


Plate 4. Extremely-fractured and vesicular Roza Member basalt from the Rosman core hole. Note the large, white plagioclase phenocrysts in the lower right of the photograph.

having particularly high concentrations of vesicles occur from: 184.0 ft. to 202 ft. bls; 223.0 ft. to 242.0 ft. bls; 259.5 ft. to 275.0 ft. bls, and; 282.9 ft. to 306.0 ft. bls.

Fracture-filling/coating materials in the Priest Rapids and Roza basalts consist of clays, iron oxides, manganese oxide, zeolite(s), and pyrite. One dark, brown-green material suggested the appearance of serpentine in hand specimen, but x-ray diffraction results indicated the material was probably a clay mineral.

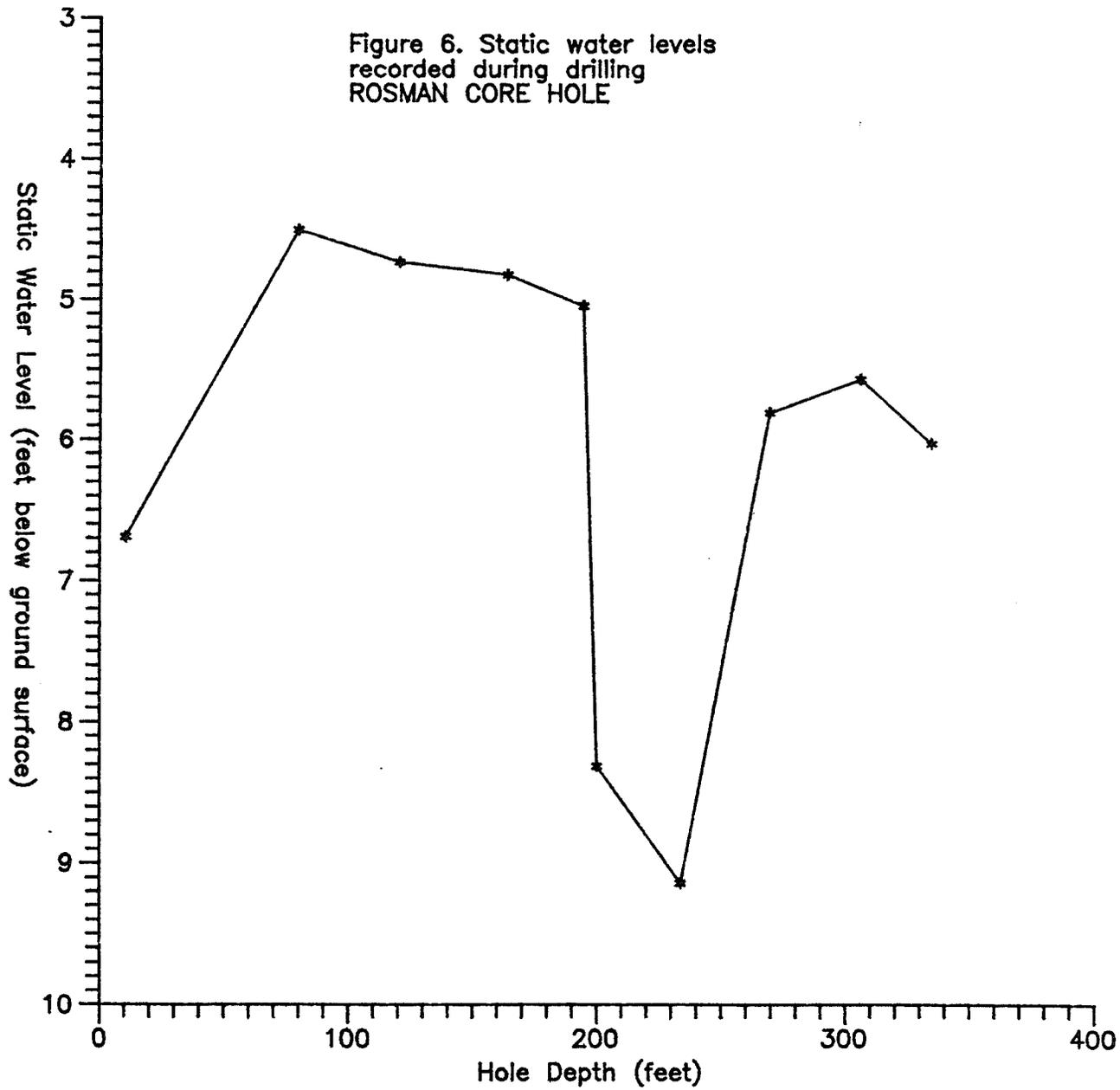
A downward hydraulic gradient is indicated by daily records of static water levels during the drilling of the Rosman core hole (Figure 6). The two outstanding readings at 200.0 ft. bls and 234.0 ft. bls are difficult to explain. It is possible that fractures near these levels were especially conductive, and may have allowed water levels to adjust overnight more quickly to ambient hydrologic conditions than during the periods of inactivity preceding the other measurements. Cuttings from subsequent drilling may have then partially plugged these fractures, thus masking their effect during later measurements. If this assumption is correct, these two lower readings may more closely reflect the composite hydraulic head at their respective hole depths than do the other readings. Supporting this assumption is the fact that static water levels in both the completed Rosman piezometers quickly stabilized at levels significantly lower than those predicted by daily readings taken during drilling.

Rosman shallow hole

During the drilling and hydrologic testing of the Rosman core hole, we decided that a monitoring point in the alluvium between the core hole and the spring would help us further resolve hydraulic interconnection of the basalt-alluvium-spring system and estimate the hydraulic conductivity of the alluvial material. The hole was drilled on July 3, 1991, to a depth of 9.5 ft. bls, and is approximately 16.5 ft. southwest of the core hole (Figure 3). Hollow-stem auger bits of 4-inch and 8-inch sizes were used to drill to total depth. Appendix A illustrates lithologies and as-built specifications for the shallow Rosman hole. Lithologies encountered in this hole are expectedly similar to the Rosman core hole, except that a thicker layer of diatomite was present (from 2.0 ft. to 4.0 ft. bls).

Static water level in the shallow hole was 5.5 ft. bls the day following drilling, and after the hole was purged of cuttings.

Figure 6. Static water levels recorded during drilling ROSMAN CORE HOLE



GEOPHYSICAL LOGGING

Washington State University, Department of Geological Engineering was contracted to complete a suite of geophysical logs for the Wagner and Rosman holes. Because of problems with caving, and a possible crooked hole, the Wagner hole was not logged. Caving in the Rosman hole also limited the depth of logging to 319 ft. bls. Functions used in the logging are; natural gamma, flowmeter (relative, vertical fluid velocity), caliper, fluid temperature, fluid resistivity, wall resistivity, and spontaneous potential. Functions using radioactive sources were unavailable. Appendix C is the geophysical log of the Rosman core hole.

Flowmeter and caliper functions exhibit marked deflections at approximately 175 ft. bls in the Rosman hole. This is nine feet above the Priest Rapids Member-Roza Member contact, as interpreted from core. Also present at approximately the same depth are large, positive and negative deflections in spontaneous potential and wall resistivity, respectively. These anomalies also approximately correspond to the flow bottom of the Priest Rapids Member, which is probably the first, significant water-producing zone encountered in the core hole. Gephart and others (1979) also found resistivity and spontaneous potential the most useful functions for identifying basalt interflow zones and water-producing horizons.

HYDROLOGIC TESTING

Injection and pumping tests were conducted at the Rosman drilling site in order to test for hydraulic connection between basalt aquifers and Baring Spring. Accurate hydraulic parameters for the formation were not sought in the initial injection test. Rather, we wished to induce a measurable rise in the pool of Baring Spring by injecting water into an isolated interval of basalt, and thereby demonstrate hydraulic connection between spring and basalt. Alternatively, we would pump the spring pool and try to observe a response in water levels in the basalt aquifer(s). We hoped that, at a minimum, a qualitative relationship might be demonstrable because of the proximity of the core hole to the spring (50 feet).

In the first test, a 24-inch-long inflatable packer was installed at the 45.5-foot-bls level in the Rosman core hole. At the time of the test, the hole depth was 195.0 ft. bls, but was back-filled with grout to 184.0 ft. bls. Water was injected into the isolated interval below the packer at a regulated, metered rate from a large tank truck using a 5-horsepower "trash" pump. The injection rate was monitored

by a "Tyme-Flyte" ultrasonic flowmeter and an impeller-type, totaling flowmeter. Plate 5 shows the equipment used in the first injection test. The levels of the spring pool and of the water in the annulus of the core hole (above the packer) were monitored during the test. Head within the isolated interval below the packer was recorded by transducer and data logger. Appendix D is the record of head (in feet) measured within the isolated interval during the test.

During 12 minutes of injection, at a rate of 15 gpm, the hydraulic head in the interval below the packer increased by almost 90 feet; from 24.4 ft. to 114.1 ft. above the level of the transducer. No response was observed in the level of the spring pool. The water level in the interval above the packer rose at a rate that was too rapid to accurately measure with equipment at hand. The latter observation indicates either a poor seal between the packer and the bore wall, or upward movement of water through fractures that intersect the bore and subtend the packer. Approximately 12 minutes into the test, fittings on the injection line leading to the packer failed suddenly and the test was abandoned. A generous fountain of water issued from the hole for several minutes after the injection line broke. Apparently, much of the water injected was not readily accepted by the fracture system in the basalt, and was ejected by force of the overpressured interstices in vicinity of the hole when the pump was turned off. We then surmised that to induce a response in the spring pool, through such a marginally-conductive medium, might require several hours and/or lower injection rates.

The following day, we pumped the pool of the spring itself for two hours at 65 gpm. During this test, water levels in the core hole (above and below the packer) and in the spring pool were recorded. We also monitored the water level in the shallow piezometer that we had installed at the end of the previous day. This piezometer is open to the alluvial, water-table aquifer and is situated roughly halfway between the spring and the core hole (Figure 3). The pumped water was discharged approximately 75 feet south of the spring pool toward the main channel of Sinking Creek. A two-minute interruption occurred 24 minutes into the test because of an unsecured oil plug on the pump engine.

Figures 7-10 show the results of the spring-pumping test. The spring pool itself was lowered 75 mm (2.95 inches) during the two hours of pumping. The pool recovered 88% of its pre-pumping level within 10 minutes of pump shutdown. Superimposed on the pumping response of the spring pool are some slight effects of intermittent withdrawals by a stock-watering pump and an ongoing, seasonal decline in the pool.

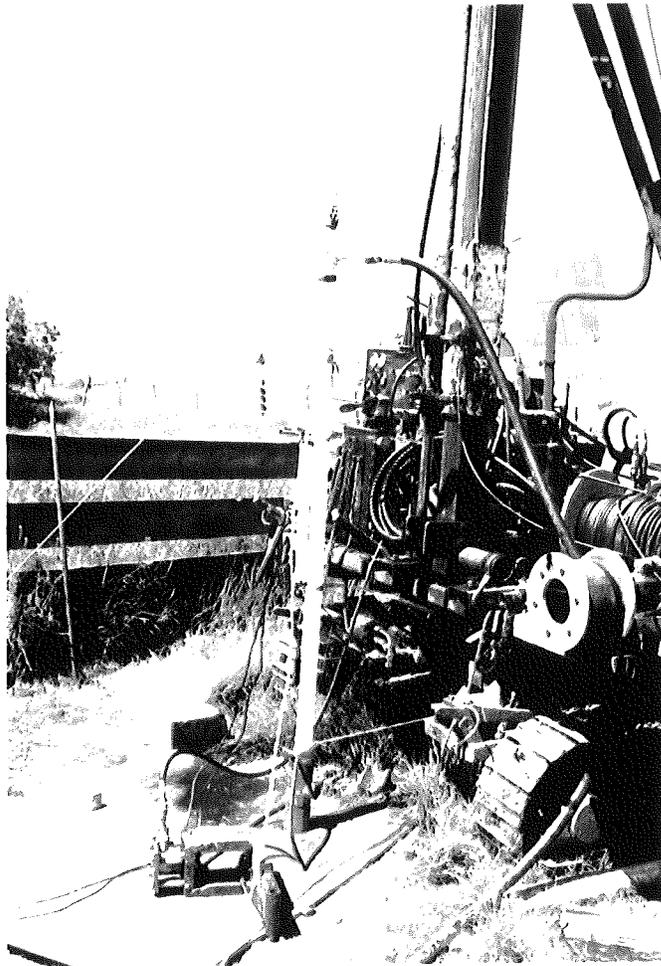


Plate 5. The apparatus used for the first injection test. Shown are the ultrasonic flowmeter (attached to the left side of the vertical injection pipe) and the adapter at the top of the pipe allowing the entry of the transducer tube.

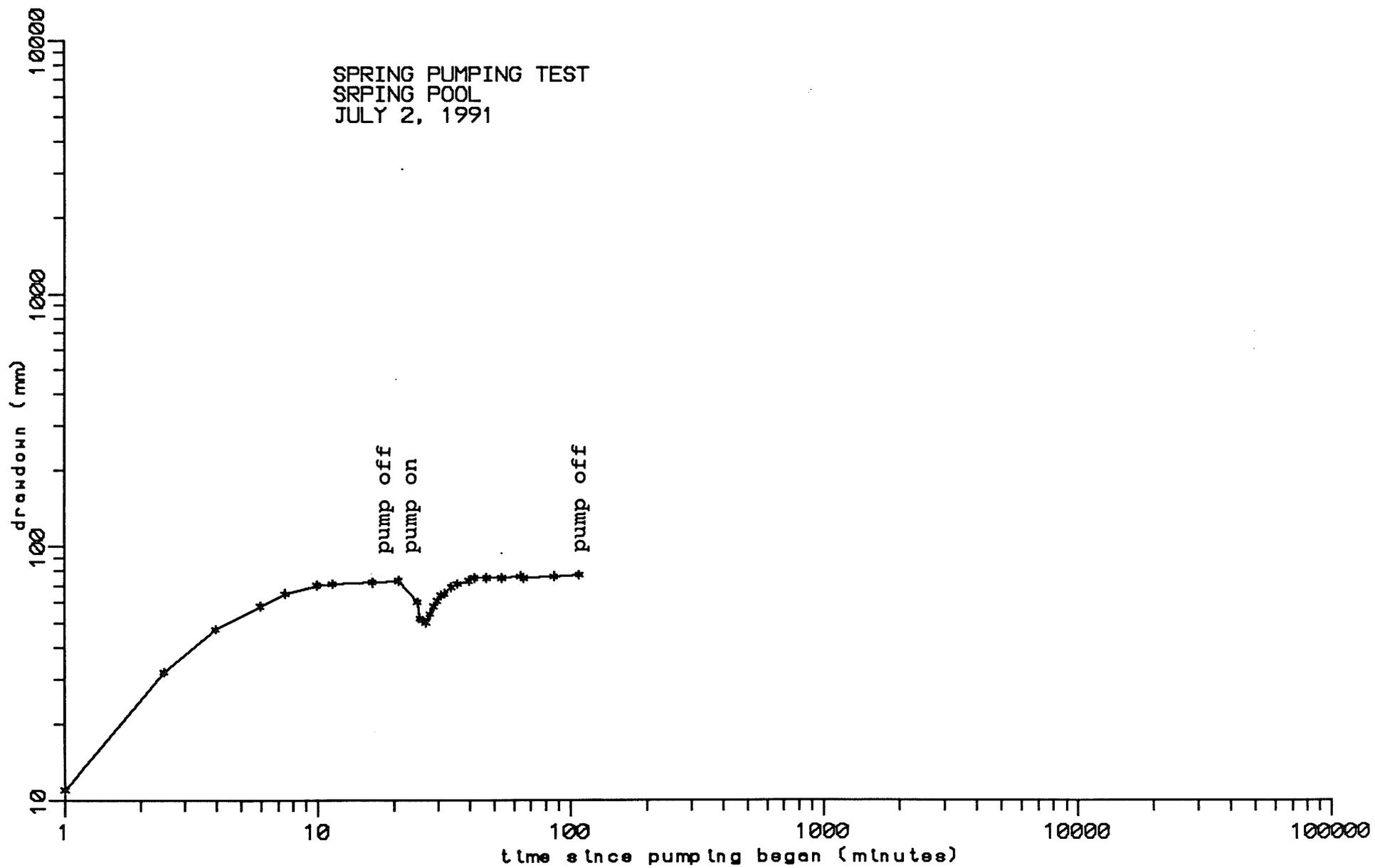


Figure 7. Time-drawdown plot of the spring pumping test--
spring pool.

SPRING PUMPING TEST
SHALLOW PIEZOMETER--31 FT. FROM SPRING CENTER
JULY 2, 1991

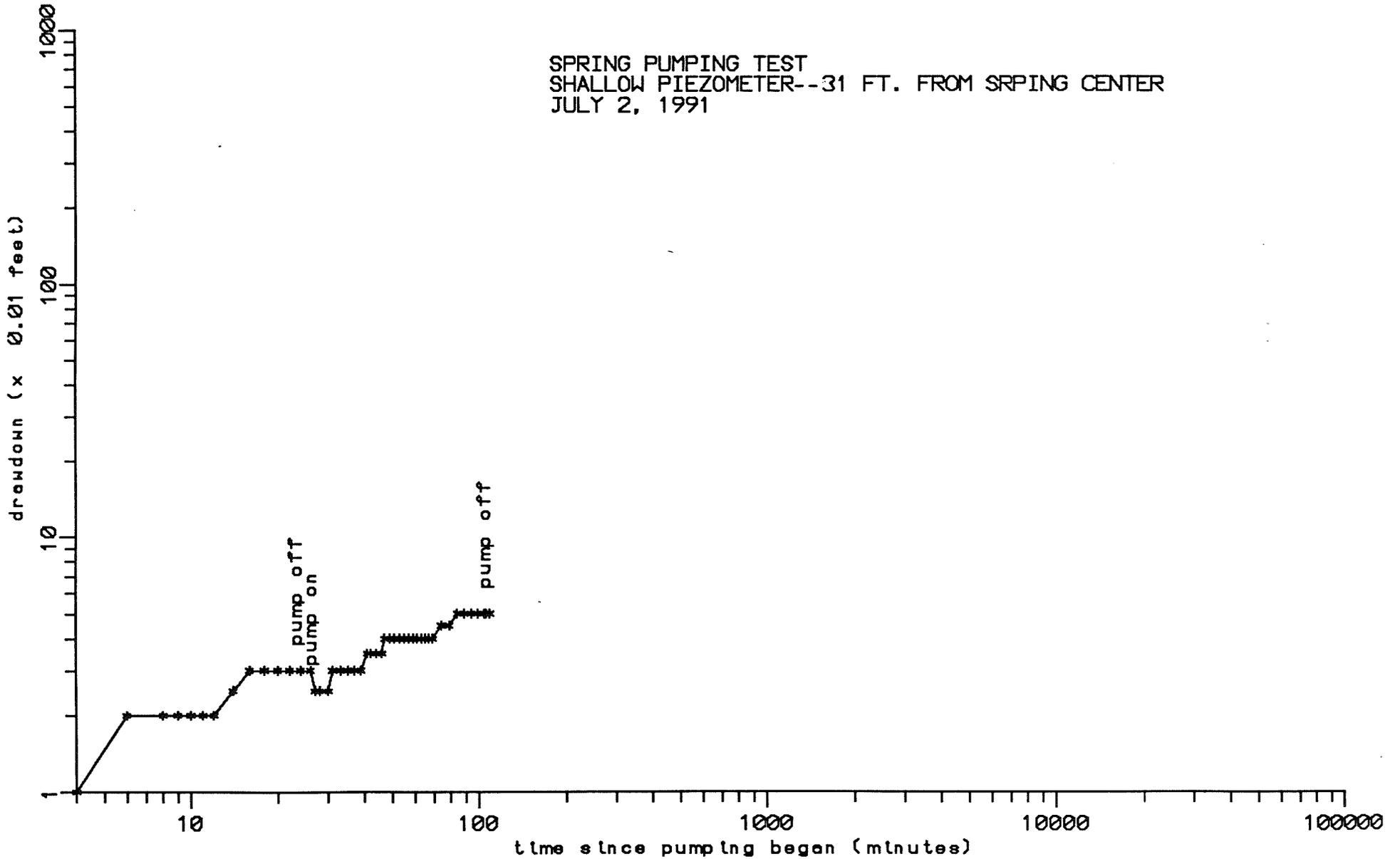


Figure 8. Time-drawdown plot of the spring pumping test--shallow piezometer.

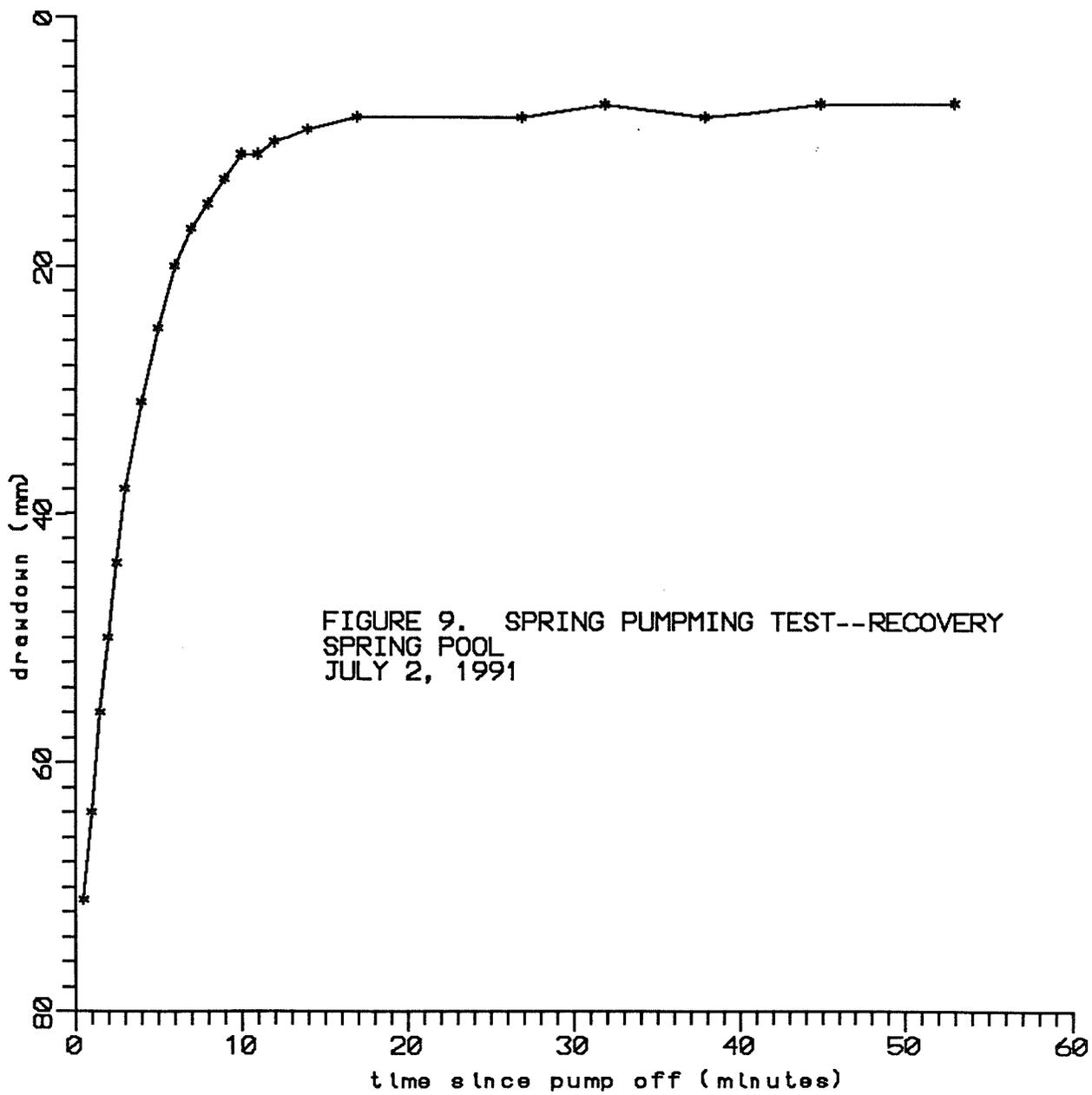


Figure 9. Spring pumping test recovery--spring pool.

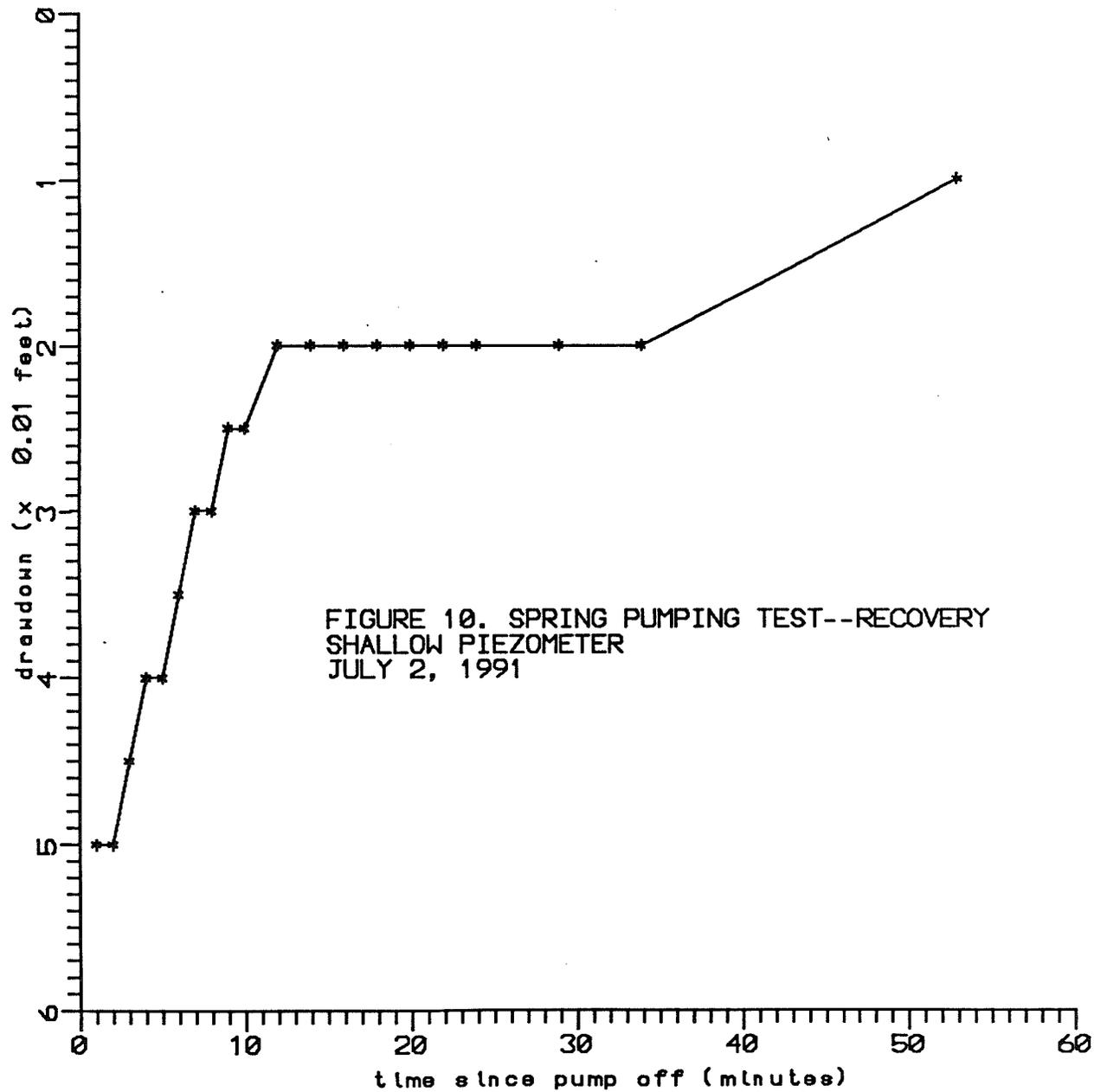


FIGURE 10. SPRING PUMPING TEST--RECOVERY
 SHALLOW PIEZOMETER
 JULY 2, 1991

Figure 10. Spring pumping test recovery--shallow piezometer.

Records of spring-pool fluctuations (Appendix E) suggest that the stock-watering pump may intermittently cause a lowering of the pool on the order of a few millimeters (few hundredths of a foot). The seasonal decline rate during the drilling project was on the order of 0.05 ft. in 24 hours.

The water level in the shallow piezometer declined 0.05 ft. during the test, and had recovered to within 0.01 ft. of static level 53 minutes after the pump was shut down. Neither of the water levels in the core hole (above or below the packer) responded to the test.

Rough analyses of the time-drawdown data in Figure 7, assuming non-equilibrium conditions, and comparison with Theis type-curves, indicate a transmissivity for the alluvium of 6000 ft.²/day, a hydraulic conductivity of 1200 ft./day, and a storage coefficient of 0.02. The value for transmissivity is probably well below the average for the alluvium over the entire coulee bottom. Shallow seismic-refraction surveys indicate that the alluvium thickens considerably toward the main channel of Sinking Creek at the center of the coulee (due south of the test area).

A final injection test was attempted in the Rosman hole on July 9, with the hole depth at 200.2 ft. bls. This time, we set the packer at 75.0 ft. bls, within a relatively fracture-free interval. Several highly fractured intervals occur between 75.0 ft. and 200.2 ft. bls. No instruments were installed in the isolated zone below the packer. The injection rate was approximately 15 gpm, as in the first test. We monitored water levels in the spring pool, the shallow piezometer, and the annular space above the packer in the core hole. At six minutes into the test, the packer assembly failed once again and the testing was abandoned for lack of time. During the six minutes of testing, no response was observed in the spring pool or the shallow piezometer.

In summary, two bootless injection tests in the Rosman core hole and a two-hour pumping test of the spring pool failed to demonstrate hydraulic connection between the Priest Rapids basalt aquifer and Baring (Rosman) Spring. These efforts were thwarted by failure of equipment during injection tests, low hydraulic conductivity of the tested intervals, and limited pump capacity. The pumping test did show that the alluvium has high hydraulic conductivity and potentially high transmissivity.

PIEZOMETER INSTALLATION

Because of the small diameter of the core holes, 1-inch, schedule 80 PVC (nominal outside diameter of 1.3 inches) was used as piezometer tubing. Tolerances, therefore, were low, in the core holes, especially in the Rosman core hole where two piezometers were installed and a 1-inch tremie was used to insert pack and sealing materials. Despite this constraint, all of the equipment and seals were installed as planned. The Rosman shallow piezometer, also 1-inch PVC, was installed in an 8-inch auger hole without difficulty. Appendix A shows as-built specifications for the four piezometers.

The injection tests had taught us that although an interval of basalt may contain numerous fractures, it may not necessarily transmit ground water. To ensure communication between water-bearing fractures and the piezometers, 20 ft. to 30 ft. of screen was installed for each piezometer within selected, densely-fractured intervals thought to be water-bearing. Twenty feet of screen was installed for the Wagner piezometer in the Priest Rapids Member basalt (from 113.7 ft. to 133.7 ft. bls), so that the bottom of the screen rests upon a bentonite-pellet seal 7 ft. above the Quincy interbed.

The two piezometers in the Rosman core hole were screened from 157.0 ft. to 177.0 ft. bls (in the Priest Rapids Member basalt), and from 285.0 ft. to 315.0 ft. bls (in the Roza Member basalt). Brief slug tests in both piezometers indicated no intra-bore hydraulic communication between the Priest Rapids Member and the Roza Member.

The shallow piezometer at the Rosman site was screened from 4.5 ft. to 9.5 ft. bls. Piezometer screens in all holes are Schedule 80 PVC, 20-slot. Sand packs were emplaced around all screens with a tremie, except in the shallow hole, where sand was poured through the hollow-stem auger bit.

Upon completion, dedicated transducers were installed in both piezometers in the Rosman core hole. At the time of completion, static water levels in the Rosman piezometers were 5.5 ft. bls in the shallow hole (measured July 3, 1991), 20.7 ft. bls in the Priest Rapids basalt (measured August 8, 1991), and 31.8 ft. bls in the Roza Member basalt (measured August 8, 1991), indicating downward flow potential at this location.

Static water level in the Wagner hole was 70.3 ft. bls a few days after the piezometer was installed.

SUMMARY

The lithologic and hydrologic information gathered at the Wagner and Rosman sites have demonstrated that: 1. The Quincy interbed is absent at the Rosman site, and thus, is not a confining horizon at this location. 2. A downward ground-water flow potential exists between the Priest Rapids basalt and the Roza Member basalt, at the Rosman site. 3. An alluvial aquifer having high hydraulic conductivity (1200 ft./day), and potentially high transmissivity, overlies the Priest Rapids basalt at the Rosman site. 4. Intense fracturing does not uniformly imply increased hydraulic conductivity in the basalt.

Confining layers at the Rosman site, especially between the Roza Member and Priest Rapids Member, are most likely the dense, flow interiors having either few fractures or fractures sealed by secondary minerals.

Two aborted injection tests in the Rosman core hole, and one pumping test of Baring Spring could not establish a direct hydraulic connection between the Priest Rapids Member basalt and the spring, nor between the surficial alluvium and the basalt. However, the results of the study do not preclude the possibility of widespread leakage between the alluvium/spring system and the basalt aquifers near the Rosman site. The pumping test did demonstrate a direct hydraulic connection between the spring and the surficial alluvium.

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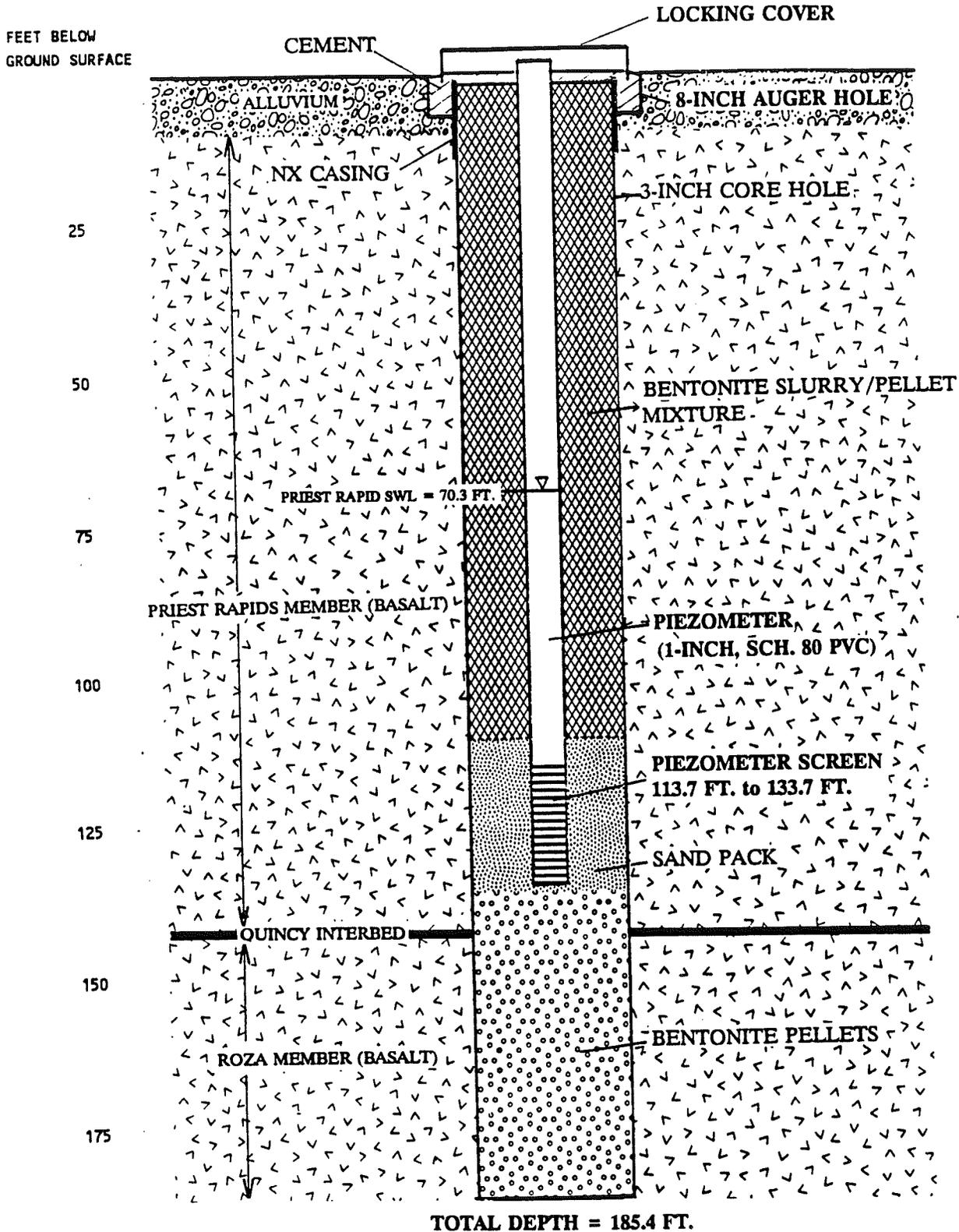
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APPENDIX A
SITE STRATIGRAPHY
AND
PIEZOMETER AS-BUILT SPECIFICATIONS

WAGNER CORE HOLE

SURFACE ELEVATION = 2090 FT. MSL

25/35-12K2



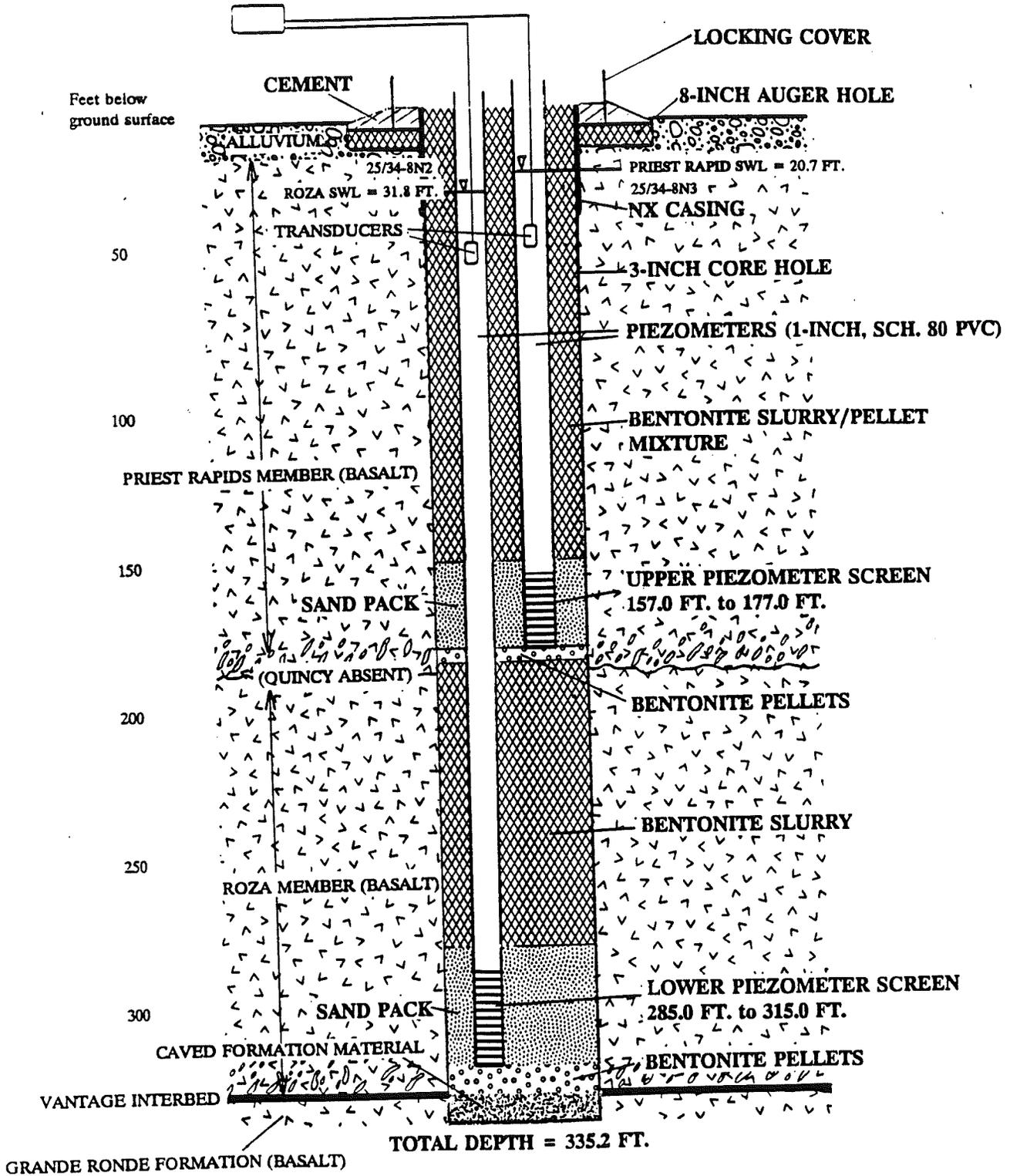
TOTAL DEPTH = 185.4 FT.

One inch equals 25 feet vertically (horizontal not to scale)

ROSMAN CORE HOLE

SURFACE ELEVATION = 2210 FT. MSL

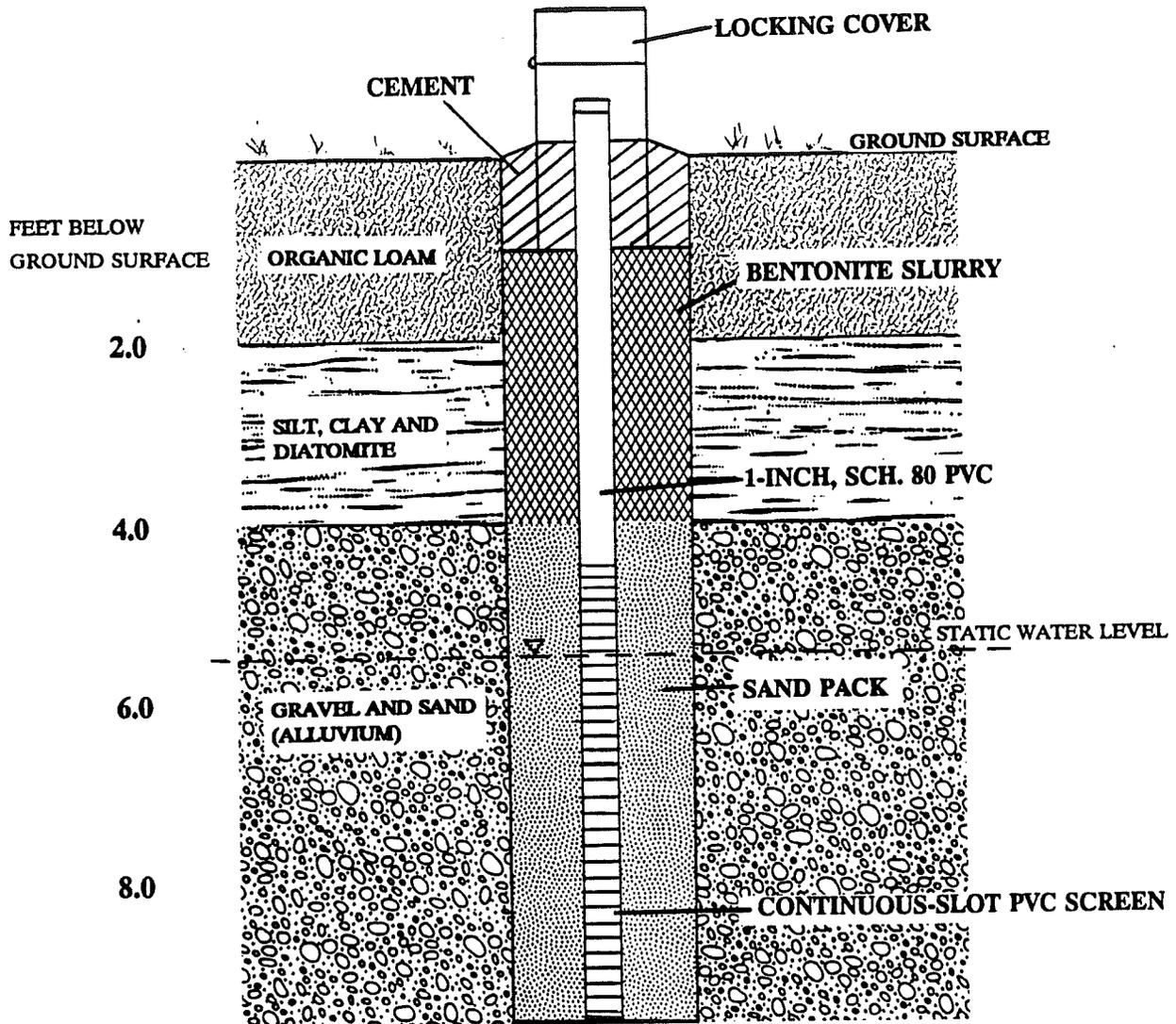
DATA LOGGERS



One inch equals 50 feet vertically (horizontal not to scale)

ROSMAN SHALLOW PIEZOMETER

25/34-8N4



TOTAL DEPTH = 9.5 FT.

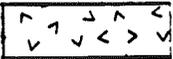
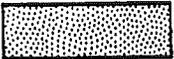
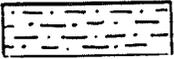
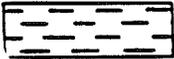
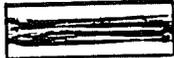
One inch equals 2 feet vertically (horizontal not to scale)

APPENDIX B

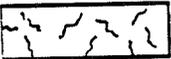
LITHOLOGIC LOGS

SYMBOLY FOR LITHOLOGIC LOGS

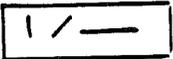
Lithologies

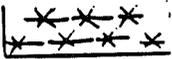
	basalt		gravel/sandy gravel
	sand/sandstone		breccia
	silt/siltstone		palagonite and similar material
	clay/claystone		volcanic glass, devit. glass

Fracturing

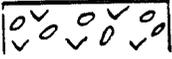
 (in graphic log)

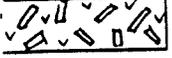
(in "Formation Condition" column)

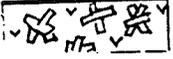
 fractures showing relative, approximate attitudes (with respect to horizontal)

 multiple orientations--pervasive

Other features

 vesicles, spherulites, and amydules (in basalt)

 phenocrysts

 glomerocrysts

DRILL LOG

HOLE NO. R - Rosman Spring ("BARING SP")

Page 1 of 14

PROJECT Sinking Creek
 CONTRACTOR Budinger's Associates
 DATE STARTED 6/21/91 COMPLETED 7/15/91
 LOGGED BY B. BARNETT

TOTAL DEPTH 335.2' LAND SURFACE ELEV. 2210 FT. MSL
 INCLINATION VERT. BEARING _____
 COORDINATES T 25N R 34E S 8, SW, SW OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Reserve Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
0.0	1/1					0		0		0' to 0.5': Sod, organic debris with, BLK, clay/silt
1.6	2/1					0				0.5' to 1.6': SILT/CLAY, PR GLN in blk, high organic content
3.1	3/1					0				
4.6	4/1					0				
5.0	5/1					0				1.6 to 3.0': Diatomite & diatomaceous earth, V. light Brn. to white. Some carbonaceous debris dissem. throughout possible volcanic ash content?
6.5						0				
10.0	6/1					0				3.0 to 3.4(?) (Poor returns): SILT, MED BRN, (Loess)
10.2	7/1					0				3.4 to 4.4(?): clay, black to dark Brn & Loess gravel. (last of drive samples)
12.0	8/1					0				4.4 to 10.2(?) (poor returns): Gravel; basalt clasts in clay/silt matrix - some sand (?) clasts are subrounded w/ low sphericity up to 15cm in long dimension
15.0						0				
17.2						0				
19.7	9/1/2					0				10.2 to 12.4(?): gravel, some sand/silt matrix (poor returns), limonite/FeO ₃ stain throughout, some surfaces (BASALT clasts)
20.0	10/2					0				
21.5						0				
23.0	11/2					0				12.4 to 25.0: BASALT, vesicular, blk, with pervasive fracturing
25.0	12/2					0				

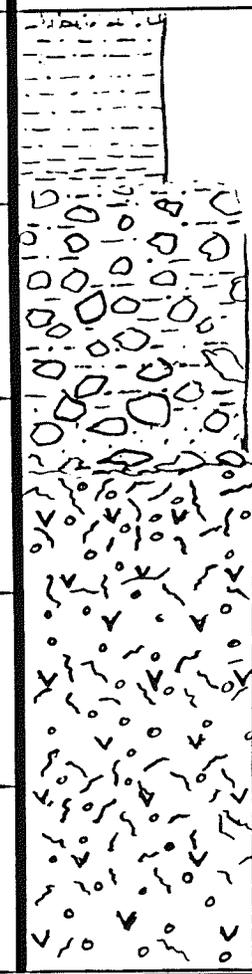
SWL @ 5.75'
 hole at 10.2'
 4/24 9:50am

lost core in hole @ -25 ft.
 lost returns @ same interval
 (some redrilled core 23.0 to 25.2)

DRIVE SAMPLES

FeO₃, Clay, Zeolite?

NX CASING



DRILL LOG

HOLE NO. Z-Rosman

PROJECT Sinking Creek
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
25.2	13/2/3	6/25 11:15 am SWL 4.56 ft BSS 20 min. after cessation of injection for drilling.		*** /		0	NX CASING	25	V V V	Started drilling @ 25.2 ft on 6/25
28.5	14/3			* / ← clay + FeOx goethite		30		30	V V V	25.0 to : BASALT, BLK, vesicular, dense, interwoven plg. (aths ~ 1mm long). vesicles get smaller become completely infilled until ~ 32.0.
33.5	15/3			*** /		0	Casing Shoe	35	V V V	- vesicles fewer @ 30.0 - 31.5 - larger vesicles 32.0 to 33.5 1 vesicle filling MnO ₂ (?) or other soft, BLK mineral
35.2	16/3			*** /		10		40	V V V	larger vesicles for short intervals (stop of drill from 45 min to 1:30 pm due to T storm)
40.2	17/4/5			*** /		50	Open NX hole	45	V V V	46.0 to ~ 60.5: - fractures mostly lack filling but w/ only occasional zeolite (?) or serpent. (?) coatings.
45.0	18/5			== (rel. open)	Serpentine (?)	60		50	V V V	

DRILL LOG

Page 3 of 14

HOLE NO. _____

PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Reserve Comments)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
50.2	19/5/6					10		50		
55.2	20/6				pyrite & zedite(?)	30	↑ NX open hole ↓	55		52.4-55.6 larger, fewer vesicles
60.2	21/6/7				large (0.75 cm) healed frzd. w/ serpent(?) or chlorite fill 60.6 to 62.2	0		60		
65.2	22/7			***	serpent. ± chlor.	0		65		
70.2	23/7/8				↑ almost no filling (open) ↓ serpent. + minor PY MnO ₂ ? MnO ₂ ?	<10		70		fewer vesicles start interval ~ 70.8
						50		75		

DRILL LOG

HOLE NO. 2 - Rosman

PROJECT Sinking Creek
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T ___ R ___ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Relative Competence) →	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
100	27/10/11			/// xxx		10		100		fewer vesicles, smaller ↓
103.5										
105	28/11/12			xxx xxx xxx	Blue invade. coating sealed w/ clay sept./clay.	20		105		incipient fractures w/ open space
110				xxx / / xxx xx xx	incipient fractures open space ↑ MnO ₂ (?) zeolite(?) ↓ bit sooty	0		110		
115	29/12/13			xxx xxx xxx xxx	incipient open space			115		- Many incipient fractures with open space. Some rotation producing "proto-breccia"
120	30/13-14		SWL @ 4.53 ft. Before start of drilling on 6/27	xxx Δ xxx xxx xxx		10		120		- stopped drilling @ ~ 121.0 on 6/26 - mismatch -
125				xxx				125		

↑ NX
open
hole
↓

DRILL LOG

HOLE NO. Z-ROSMAN Sp.

Page 7 of 14

PROJECT Sinking Creek
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T R S OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1"=5' Basic Geology: rock types, structures, alteration. (Reserve Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
150.9	34/16	150.9 - Jammed core barrel			Pyrite 150.9 to ~154.0	2	open NX hole	150		V. SMALL VESICLES throughout, but concentr. in some short intervals - very subtle. ~157° driller reports light brn cuttings in return flow, prob fract. fill material. drilling rate has slowed somewhat - Pyrite in fract. mostly pseudo-morphic after Fe ₃ O ₄ - occurs in "spots" on fract. surfaces (photo @ 168.0) - Stopped drilling at 164.7 on 6/27 ~164° ft - driller noted metal shavings in returns - could be bit deterioration. turned out to be pyrite from fract. severely fractured @ 172.5 to ~175.5 - friable, coating of FeOx (?) paleo-water table? - LACK of returns
	35/16-17			XXXXX XXXXX XXXXX ↓ some part heated		0		155		
154.8	36/17	Jamming core barrel		XXXXX		30		160		
157.8	37/17-18			Pyrite ↑ v little fill - some grn-blue coatings		20		165		
164.7	38/18-19			Pyrite ↓				170		
171.7	39/19	Poor recovery - barrel jammed		XXXXX XXXXX XXXXX XXXXX Red-brn (FeOx?) + clay		0		175		

DRILL LOG

HOLE NO. Z-Rosman Sp.

PROJECT Sinking Creek

CONTRACTOR _____

DATE STARTED _____

COMPLETED _____

LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____

INCLINATION _____ BEARING _____

COORDINATES T _____ R _____ S _____ OTHER _____

SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale <u>1" = 5'</u> Basic Geology: rock types, structures, alteration. (Release Competency)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
175						0		175		
175.9				XXXX	Red-Brn clay - small apertures	0				- driller reports lack of returns except at end of runs
177.3	40/19	Poor recovery	62 gravel gm		serpentin/clay? w/ noiled	0				
180	41/19-20				thick (1-1.5cm) grey-white clay	10				180° Vesicles - some layers of larger vesicles. at angle of ~45°
185					red-brn clay in fract.					red-brn clay fract. minor calcite, compressed vesicle layer (photo) Chalcedony - no Quincy - ?
185.2					open + serped. clay					extremely fractured, friable ROZA
185.2	42/20-21				clay - blue-grn. 0.5cm					184° very vesicular MBR at ~184°
190					Blue - halloy? quartz + Fe Mn Ox					Some vesicles filled w/ clay fracture-filling material near fract. Some will allow pass of air through core
190					Serp./clay blue/blk + zeol.	10cm				184.0 to 225°: BASALT, porphyritic w/ Plag. up to 1cm long.
190					MgO ₂ + clay + blue coatings	grn				larger vesicles progressively downward, some coated w/ MnO ₂ + blue/grn mineral - some partially im-filled w/ same - small (<1mm) vesicles throughout
195					clay, zeolites	10				6/20 Stopped drilling at 195° to cement hole for injection test. found that cement was up to 184° ft, but after redrilling only 2 small bridges were present
195.2	43/21	started here on 7/8 SWL @ 5.44 BGS inside PACKER - 5.18 BGS in annulus.			zeolites MnO ₂	30				stopped drilling again at 200.2' to test interflow zone 7/8.
200								200		

NX open hole

DRILL LOG

HOLE NO. Z-Rosman

PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1"=5' Basic Geology: rock types, structures, alteration. (Reserve Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
225	47/24-25	- driller reports blocking	-	444 BLK clay green/white clay - some open	CLAY	0	↑ Open NX hole ↓	225		vesicles at base of glass horizon preferent. aligned w/ contact 225.4 to 225.8 - clay, black (Devitr. glass?) fracture filling? Resinous luster on fresh surface. Some BLK-grn glass still present. Both glass and "BLK clay" display conchoidal fract. pattern. 225.8 - 229.0 Vesicular porphyritic BASALT - densely fractured 229.0 - 229.2: clay, grn-brn to blk, devitr. glass 229.2 to 236.4: BASALT, vesicular, porphyritic, intensely fractured-fragmented. 236.4 to 243.3 BASALT, vesicular, not as fractured/fragmented as overlying interval. Vesicles become larger but fewer with depth this interval 243.3 to 259.5: BASALT, PORPHYRITIC med-dark gray w/ flag. thens up to 1cm. some vesicles partly clay-filled (photo)
232.1	48/25	stopped here on 7/9	scraped + grn/brn to white clay	0		230				
233.9	49/25-26	SWL in hole at 9.14 @ 8:58 am driller reports more consistent returns - pump @ 12.7 gpm	red/brn clay > 3cm	10		235				
240.7	50/26		mostly healed	0		240				
244.4	51			20		245				
							250			

DRILL LOG

HOLE NO. 2 - Rosman Sp.

Page 11 of 14

PROJECT Sinking Creek
 CONTRACTOR Budinger
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
250 250.4	52/27	short runs due to blockage/fracts.		—		40	↑ Open NX hole ↓	250		259.5 to 264.3 BASALT, very vesicular, possibly scoriaceous, extremely fractured, fragmented. Vesicles gradually diminishing downward beyond this interval (porphyritic). large vesicles pipe vesicle (5 cm in length) 259.0 to 264.3: Vesicular basalt, very fractured/fragmented 264.3 to 290.8: BASALT, Vesicular, porphyritic 264.5 - some vesicles wholly or partly clay-filled. Large & small vesicles - seems like two distinct populations/modes - vesicle "trains" - vertical bands of vesicles larger than most (very porous - after drying exterior interior of core retained moisture, possibly minor permeability)
254.1 255	53/27-28			← mostly clay 2mm ← healed ← open(?) + MnO ₂ Serp.	70	255				
260 260.2	54/28			XXXX light gray/brown clay XXXX mostly open(?) XXXX	0	260				
264.1 265	55/28-29			—	50	265				
270 270.2	56/29-30	Stopped here on 7/10 SWL on 7/11 5.80 BPS @ 5AF 2H		— — — — —	very few fracts mostly horiz. MnO ₂ + open	50		270		
275							275			

DRILL LOG

HOLE NO. Z-Rosman

Page 13 of 14

PROJECT Sinking Creek
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNEST

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1"=5' Basic Geology rock types, structures, alteration. (Reserve Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
300	64/31	Stopped at 301.5 on 7/11		XXX	fractures filled w/ grn/gry clay	0		300		
301.5		No SWL measurement		XX	rubble					
	65/31	Drilling faster less backspin			clay-filled spiracle?	0				
305		Driller added polydrill to help w/ cuttings (floc out)		XXX		0		305		
306	66/32	polymer (mistake on first run - 1 gal to trip out of hole)		XXX	pyrite	0				
307	67/32	SWL on 7/15 = 5.5 gpm 10:08 am		XXX	open black glass like	10				
310	68/32			XXX		0				
310.5	69/32			XXX		0				
311.7	70/32	poor recovery (~50%)		XXX		0				
312.2				XXX		0				
315	71/32	return flow light grey-grn.		XXX	light blue-grn. coating + blk glass like	0				
316				XXX		0				
320	72/33			XXX	open + blue clay coating	0				
				XXX		0				
				XXX		0				
				XXX	much pyrite	0				
325				XXX		0		325		

DRILL LOG

HOLE NO. 2-Rosman

Page 14 of 14

PROJECT SINKING creek
 CONTRACTOR _____
 DATE STARTED 6/21/91 COMPLETED 7/15/91
 LOGGED BY BARNETT

TOTAL DEPTH 335.2' LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1"=5' Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
325.2								325		
330	73/33-34			— — —	3cm. Horizont. Fract filled w/ hzrd dk grn. clay	30		330		325.2 to 325.5? clay dark gray-grn. VANTAGE HORIZON Bright gray-green returns ~ 320 to 330 - some loss of recovery ~ 325-330? Some sspilitic clay at ~ 325.4 - 325.6. - weathered basalt. some vesicles clay-filled
335				/	mostly open - no fill.		↑ open NX hole	335		325.5 to 335.2: - BASALT, MED. gray, vesicular, aphyric, w/ <1mm plaq. ls the densely interwoven. GRANDE RONDE FM. Stopped drilling on 7/15 @ 335.2 - T.D.
335.2	T.D.			XXX				335.2 = T.D.		
340								340		
345								345		
350								350		

SWL on 7/16
 @ 302m = 98.66 ft BGS.

100
 200
 300
 400
 500
 600
 700
 800
 900
 1000
 1100
 1200
 1300
 1400
 1500
 1600
 1700
 1800
 1900
 2000
 2100
 2200
 2300
 2400
 2500
 2600
 2700
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 8700
 8800
 8900
 9000
 9100
 9200
 9300
 9400
 9500
 9600
 9700
 9800
 9900
 10000

1.2 from 153-180
2.7

DRILL LOG

Page 1 of 8

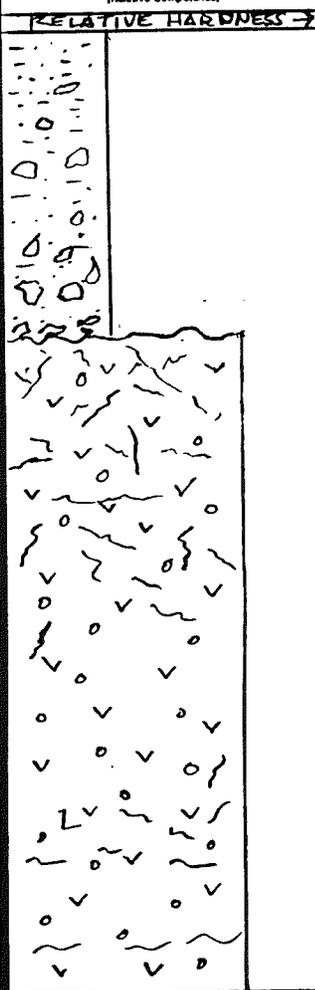
HOLE NO. 1-WAGNER SP.

PROJECT SINKING CREEK
 CONTRACTOR BUDINGER & ASSOC.
 DATE STARTED 6-11-91 COMPLETED 6-19-91
 LOGGED BY BARNETT/covert

TOTAL DEPTH 185.4 FT. LAND SURFACE ELEV. 2090 FT. MSL
 INCLINATION VERTICAL BEARING _____
 COORDINATES T25N R 32E S SE, SE 12 OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule Hole Construction	Footage	Graphic Log Scale 1" = 5 FT. Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
0								0		0.0 to 8.0: Loess silt, mixed with basalt frag./gravel some cobble-sized, med. Brn
5		dry					8" AWG. 8" AWG. HQ CASING	5		
10	1/1 2/1 3/1 4/1 5/1 6/1			(dips) 40°, 10° (70°)		0 0 0 0 0	SHOE	10		8.0 to 130.6: BASALT, DARK BRN to BLK, Aphyric, Amygdoles of clay min. (?) + FeOx evenly dist. throughout but only <1% of volume. Amyd. ~ 1-2 mm, gives rock a "speckled" appearance (Priest Rapids mbr.) (?) Densely fractured to ~ 16.0 ft. Densely fractured in some intervals throughout
15	7/1					0	NX hole	15		
15.3										
18.0	8/1					44		18.0		
20	9/2					60		20		
20.4										
22.8	9/2					0		22.8		
25	9/2					0		25		

Driller reports water levels fluctuating between 13-15 ft. throughout day on 6-14-91



DRILL LOG

HOLE NO. 1 - WAGNER Spring

Page 2 of 8

PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT / covert

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale $1" = 5'$ Basic Geology: rock types, structures, alteration. (Relative Compaction)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
25	9/2/7			///		10		25		
30	10/2/3			vert predom.	FeOx, clay, minor zeolite	0	NK open hole	30		Fracturing ~26.5 - 30
35	11/3			—	Chlorite (serpent?) zeolite calcite?	0		35		Dense fracturing ~31.5 - 33.0
40	12/3			—		0	NK open hole	40		Started drilling ~ 10:30 on 6/17 hole at 42.7 ft. water injected into hole. infiltrates quickly - drilling about 15 ft./hr.
43.4	13/3			—		0		43.4		
45	14/3/4	dry E-tape deteds no water at 10:30 am on 6/17/91 - hole at 42.7 ft.		—		<10		45		
48.4	15/4			—		~20		48.4		
50	16/4	← start on 6/17		—		0		50		DENSE FRACTURING 42' → 44' Semi VERTICAL
50	17/5			—	MnO ₂ FeOx	40		50		
50	18/5/6			—				50		

DRILL LOG

HOLE NO. 1 - WAGNER SP.

Page 3 of 8

PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT / Covert

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale <u>1" = 5'</u> Basic Geology: rock types, structures, alteration. (Relative Compaction)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
59.0	18/5/6					80		50		Photo of "Halloysite" @ 54.9'
	19/6			↑	Blue, clay like (Halloysite?) esp. 54-57 some chlor. or serpent. some int.	70		55		
	20/6/7			↓	chlorite / FeOx	40	open N1 hole	60		Photo of "healed" fractures w/ clay/chlor. @ 66.6' w 1-2 mm apertures max.
63.6	21/7			//		60		65		
68.7	22/7	- inner tube blocked -			/			70		
74.2	23/7/8			-				75		
	24/8			-						

DRILL LOG

HOLE NO. 1 - WAGNER SP.

Page 5 of 8

PROJECT Sinking Cr.
 CONTRACTOR _____
 DATE STARTED 1 COMPLETED _____
 LOGGED BY BARNETT / covert

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale <u>1"=5'</u> Basic Geology: rock types, structures, alteration. (Requires Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
100								100	✓ ✓ ✓	Drilling rate approx. 12 ft./hr. Recovery ~ 100%
	31/11			—		80			✓ ✓ ✓	
105.4				fract. ^{see} every 0.8'		80	open NK hole	105	✓ ✓ ✓	
	32/11/12			φ 2 ~55° most common		80			✓ ✓ ✓	
110.4				Pyrite 2 bands in fract. 91.0 to 111.1		80		110	✓ ✓ ✓	
	33/12			Chlor./ serpent.		80			✓ ✓ ✓	
116.4				↓		50		115	✓ ✓ ✓	
	34/12/13			↓		50			✓ ✓ ✓	
120.7 122.7	35/13	Jammed core barrel		↓		0		120	✓ ✓ ✓	
	36/13			↓		60			✓ ✓ ✓	
125				↓		60		125	✓ ✓ ✓	

~ 116° to 122°: BASALT CONTAINS
 remnant phenocrysts ~ 1mm
 (olivine → chlorophanite?)
 or idding site
 - mostly Hazled fract 117.9 - 118.9

- stopped drilling ~ 300
 pm
 At 120.7 ft. started ~ 9:30 am on
 6/19. suc = 38.29 ft.
 at 123° Vesicles/Amygdloles
 reappear

DRILL LOG

HOLE NO. 1- WAGNER SP

Page 6 of 8

PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT / Covert

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale 1" = 5' Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
125 125.4	37/13/14			large vesicles		50		125	relative HARDNESS →	large vesicles coated with blk. botryoidal mineral w/ light brn streak (flow bottom (pinstriped))
130 130.4	38/14					0	NK open hole	130		130.6 to 133.9: Scoriaeous BASALT and flow breccia vesicles filled w/ limonite & other Fe OXs (Photo)
135 134.2	39			chlor./serpent other blk min.		30		135		ABUND. FRACTS 135.6 to 137.0 - Bx + vesicles - drilling rate ~ 20 ft./hr. vesicular 139.8 to 140.4
140 140.4	40					0		140		141.7 to 142.4: claystone, med. grn-brn, possibly good clay and/or glass at base of flow. Quincy "interbed"?
145 144.8	41/16	JAMMED	inner tube			0		145		142.4 to 145.4 BASALT, vesicular (Flowtop) w/ plagioclase esp. in vesicles (photos) (BEEIN POZA MBR.) Also Halloysite(?) Bright blue-grn coating on vesicles
150 145.4	42/16					40		150		

DRILL LOG

HOLE NO. 1-WAGNER SP

Page 7 of 8

PROJECT SINKING CREEK
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT/COVERT

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale <u>1" = 5'</u> Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
150	43 16/17			/		20		150		<p>(cont'd) locally vesicles are smaller but more concentrated (eg. 155.5 and 162.5) Plagioclase pheno. become more prominent downward; up to 1.5-cm-long plag. laths. Aphyric groundmass.</p> <p style="text-align: center;">↑ (ROZA FLOW TOP) ↓</p> <p>165.0 ~ 178.0: vesicles larger, fewer - some vesicles > 3.5 cm</p> <p>- LARGER VESICLES</p>
155 155.4			/		30	NX open hole ↑ ↓	155			
160 160.4	44/17		/		10		160			
165 165.4	45/17/18		/		50		165			
170 170.4	46/18		/		50		170			
175	47/18/19		/				175			

DRILL LOG

HOLE NO. 1-WAGNER SP

Page 8 of 8

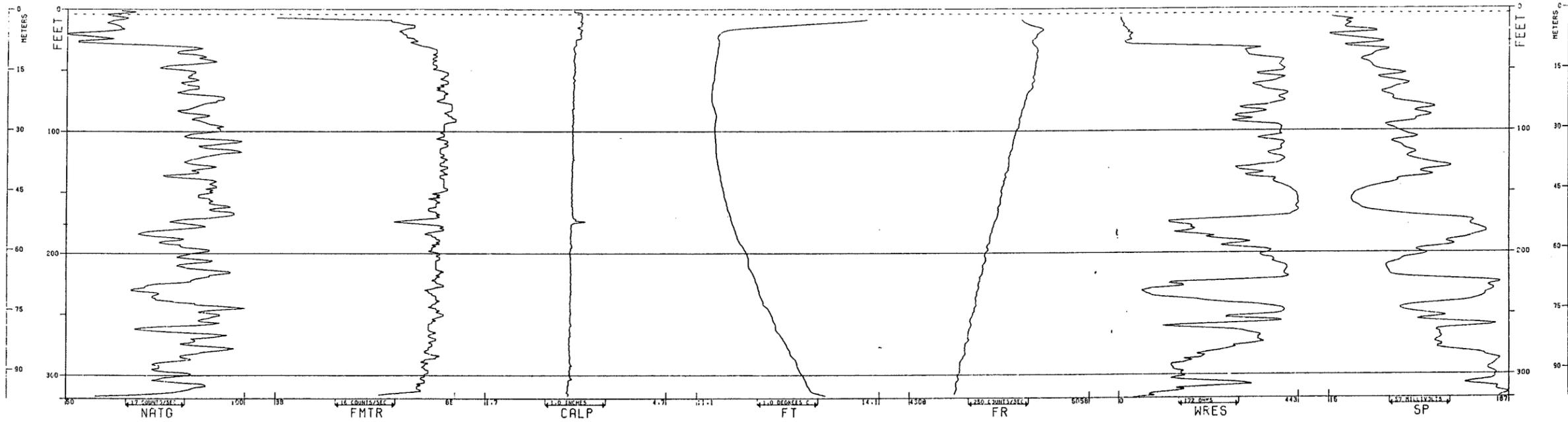
PROJECT _____
 CONTRACTOR _____
 DATE STARTED _____ COMPLETED _____
 LOGGED BY BARNETT / covert

TOTAL DEPTH _____ LAND SURFACE ELEV. _____
 INCLINATION _____ BEARING _____
 COORDINATES T _____ R _____ S _____ OTHER _____
 SURVEY REFERENCES _____

Footage	Hole Condition			Formation Condition			Casing Schedule	Footage	Graphic Log Scale <u>1" = 5'</u> Basic Geology: rock types, structures, alteration. (Relative Competence)	Lithologic Descriptions, Drilling Notes
	Run/Box#	Water (gpm)	Pit Loss (gpm)	Fracture Orientation	Fracture Filling	RQD (%)				
175 175.3	48/19			/			↑ Open NX hole ↓	175		SMALLER VESICLES - Stopped drilling @ 185.4' on 6/19 - SWL @ 79.98 on 6/20 10:30 AM Regan preparing to move off site
182.0	49/19/20			/		180				
185.4				/		185				
								190		
								195		
200								200		

APPENDIX C

GEOPHYSICAL LOG OF THE ROSMAN CORE HOLE



WASHINGTON STATE UNIVERSITY
COLLEGE OF ENGINEERING
 GEOLOGICAL ENGINEERING SECTION
 WELL LOG PROCESSING SYSTEM

NAME OF WELL	ROSMAN
DATE LOGGED	07/17/91
STATE	WASHINGTON
COUNTY	LINCOLN
LOCATION	25N/34E-08N1
SURFACE ELEVATION	2207
TOTAL DEPTH LOGGED	0-319
DEPTH TO WATER LEVEL	4
CASING & LINERS (P=PVC):	
0- 30= 3	

LEGEND

LOG TITLES

- NATG - NATURAL GAMMA
- GG - GAMMA GRAMM
- HN - NEUTRON NEUTRON
- NO - NEUTRON OPHM
- FT - FLUID TEMPERATURE
- FR - FLUID RESISTIVITY
- FA - FLUID RESISTIVITY
- CALP - CALIPER
- SP - SPONTANEOUS POTENTIAL
- WRES - WALL RESISTIVITY
- FMTR - FLOW METER, 1" DOWN
- SV - SATE VELOC
- SA - SATE AMPLITUDE
- SW - SATE WAVEFORM
- SN - SATE NORMAL
- LN - LOG NORMAL
- CONS - CONSTRUCTION
- * TRANSMI-RECEIVE OSP (0-10")
- * P.S.T WAVES, CASING

BELOW WATER LEVEL:

- DENSITY (OO-LOG) INCREASES - - - - -
- POROSITY (NN-LOG) INCREASES - - - - -
- NOTE: SCALE MAY CHANGE ABOVE WATER LEVEL

CONSTRUCTION SYMBOLS

FORMATION	----
GRAVEL	
GRUIT	----
BENTONITE	
SCREEN	----

07/22/91 KRS

ROSMAN CORE HOLE--BELOW PACKER
FIRST INJECTION TEST
JULY 2, 1991

Terrasience Systems Ltd., Terra8 Series Data Printout

Terra8 Data Collection Report

Firmware Version: 0
Time of Data Dump: 07/02 12:24:52.54
User Comment: Injection test of Priest Rapids basalt. Rosman corehole
Default Scan Interval: 00/00 00:00:30.00
Number of Logs to Memory: 122
Remaining Data Memory: 22379
Filename: c:\sinkcrk\drilling\rosman.wel\prelim1.oat

Terra8 Channel Setup:

Number of Declared Analog Channels = 1

Ch#	Description	Units	Delay	M	B
1	50 psi transducer	ft	100	22.550	1.910

Terra8 Channel Setup:

Number of Declared Digital Channels = 0

Ch#	Description	Units	Delay	M	B
-----	-------------	-------	-------	---	---

APPENDIX D

DATA FROM FIRST INJECTION TEST

Time	Analog#01 - Feet of water above transducer
07/02 10:03:35.90	7.547
07/02 10:04:07.37	25.024
07/02 10:04:09.39	25.024
07/02 10:07:25.58	44.755
07/02 10:07:29.87	44.755
07/02 10:08:42.04	64.712
07/02 10:49:07.67	24.460
07/02 10:49:12.25	24.573
07/02 10:49:16.19	24.460
07/02 10:49:18.76	24.460
07/02 11:22:36.58	24.234
07/02 11:22:41.20	24.460
07/02 11:22:57.77	24.347
07/02 11:23:27.34	24.347
07/02 11:23:57.34	24.347
07/02 11:24:27.34	24.460
07/02 11:24:57.34	24.460
07/02 11:25:27.39	24.347
07/02 11:25:57.34	24.347
07/02 11:26:27.34	24.460
07/02 11:26:57.34	24.460
07/02 11:27:27.34	24.460
07/02 11:27:57.34	24.347
07/02 11:28:27.39	24.234
07/02 11:28:57.34	24.460
07/02 11:29:27.34	24.347
07/02 11:29:57.34	24.347
07/02 11:30:27.34	24.460
07/02 11:30:57.34	24.347
07/02 11:31:27.39	24.460
07/02 11:31:57.34	24.460
07/02 11:32:27.34	24.347
07/02 11:32:57.34	24.347
07/02 11:33:27.34	24.460
07/02 11:33:57.34	24.347
07/02 11:34:27.39	24.347
07/02 11:34:57.34	24.460
07/02 11:35:27.34	24.347
07/02 11:35:57.34	24.347
07/02 11:36:27.34	24.460
07/02 11:36:57.34	24.460
07/02 11:37:27.39	24.460
07/02 11:37:57.34	24.347
07/02 11:38:27.34	24.460
07/02 11:38:57.34	24.460
07/02 11:39:27.34	24.460
07/02 11:39:57.34	24.347
07/02 11:40:27.39	24.347
07/02 11:40:57.34	24.347

Time	Analog#01
07/02 11:41:27.34	24.347
07/02 11:41:57.34	24.460
07/02 11:42:27.34	24.347
07/02 11:42:57.34	24.347
07/02 11:43:27.39	24.347
07/02 11:43:57.34	24.460
07/02 11:44:27.34	24.347
07/02 11:44:57.34	24.460
07/02 11:45:27.34	24.347
07/02 11:45:57.34	24.347
07/02 11:46:27.39	24.347
07/02 11:46:57.34	24.460
07/02 11:47:27.34	24.460
07/02 11:47:57.34	24.460
07/02 11:48:27.34	24.460
07/02 11:48:57.34	24.234
07/02 11:49:27.39	24.234
07/02 11:49:57.34	24.460
07/02 11:50:27.34	24.460
07/02 11:50:57.34	24.460
07/02 11:51:27.34	24.460
07/02 11:51:57.34	24.460
07/02 11:52:27.39	24.347
07/02 11:52:57.34	24.347
07/02 11:55:14.01	24.347
07/02 11:55:43.34	24.347
07/02 11:56:13.34	24.234
07/02 11:56:43.55	24.347
07/02 11:57:01.30	37.764
07/02 11:57:07.31	90.644
07/02 11:57:12.32	83.879
07/02 11:57:17.31	87.938
07/02 11:57:22.31	89.404
07/02 11:57:27.32	87.487
07/02 11:57:32.31	88.727
07/02 11:57:37.31	89.968
07/02 11:57:42.32	90.757
07/02 11:57:47.31	92.448
07/02 11:57:52.32	92.899
07/02 11:58:02.31	96.056
07/02 11:58:12.31	98.198
07/02 11:58:22.31	101.581
07/02 11:58:32.31	102.708
07/02 11:58:42.31	104.512
07/02 11:58:52.31	103.723
07/02 11:59:02.31	103.949
07/02 11:59:12.31	105.753
07/02 11:59:22.31	105.865
07/02 11:59:42.31	109.248
07/02 12:00:02.30	108.008

BEGIN INJECTION TEST

Time	Analog#01
07/02 12:00:22.30	109.135
07/02 12:00:42.31	110.037
07/02 12:01:02.31	108.684
07/02 12:01:22.31	110.714
07/02 12:01:42.31	109.473
07/02 12:02:02.31	109.361
07/02 12:02:22.31	111.165
07/02 12:03:02.31	111.954
07/02 12:03:42.31	111.841
07/02 12:04:22.31	112.292
07/02 12:05:02.31	111.165
07/02 12:05:42.31	112.969
07/02 12:06:22.31	112.292
07/02 12:07:02.31	113.983
07/02 12:07:42.31	114.885
07/02 12:08:22.31	114.096
07/02 12:09:22.31	16.116
07/02 12:10:22.31	15.553
07/02 12:11:22.31	15.214
07/02 12:12:22.31	14.876
07/02 12:13:22.31	14.651
07/02 12:14:22.31	14.538
07/02 12:15:22.31	15.553

INJECTION SYSTEM FAILURE

APPENDIX E

RECORD OF BARING ("ROSMAN") SPRING WATER LEVELS

BARING ("ROSMAN") SPRING (25N/34E-8N1s) WATER LEVELS

DATE	TIME	WATER LEVEL (MM below datum)	COMMENTS
JUNE 24	10:00	-----	staff gage reading (in feet) = 6.91
24	12:00	-----	staff gage reading (in feet) = 6.89
24	15:00	-----	staff gage reading (in feet) = 6.90
25	9:30	-----	staff gage reading (in feet) = 6.86
25	11:00	00154	withdrawing water for drilling.
25	11:51	00153	
25	pm?	-----	instrument reset
26	9:15	00152	
26	14:33	00155	
26	14:48	00156	
26	15:48	00150	pumping from pool for drilling had ceased about 15:38
26	16:23	00149	
27	8:34	00156	
27	10:07	00157	
27	10:42	00162	pumping from pool for drilling
27	11:45	00165	"
27	12:46	00469	instrument reset (00165 - 00469)
27	15:06	00470	
27	16:18	00472	just stopped drilling
28	8:30	00477	before drilling started
28	13:23	00480	drilling for 3 hrs.--pump running

BARING ("ROSMAN") SPRING (25N/34E-8N1s) WATER LEVELS

JUNE 28	15:20	00476	stopped drilling 1 hr. ago--T-storm in vicinity
JULY 1	9:20	00500	staff gage = 6.67
2	9:15	00510	staff gage = 6.63
2	15:00	00514	injecting water into core hole over last 3 hrs.
2	16:55	00513	no activity for 2 hrs.
3	9:36	00522	pump on intermittently
8	9:57	00589	no activity
8	12:28	00591	stock pump on intermitt.
8	17:25	00593	no activity
9	9:09	00607	stock pump running
9	10:28	00607	6 minutes into injection test
9	11:04	00608	35 min. after test ended
9	14:19	00604	drilling since 11:30
9	16:08	00604	drilling stopped 1/2 hr ago
10	9:02	00615	no activity
10	11:20	00617	stock pump running
10	17:17	00619	drilling stopped >1 hr. ago
11	9:53	00629	no activity
11	13:46	00637	drilling since 10:45
11	16:06	00640	drilling stopped 1\2 hr. ago--stock pump running
12	9:47	00650	drilling water faucet left on all previous night
15	9:35	00689	no activity--storm moved in at 9:40
15	12:38	00697	drilling--withdrawing water since 10:30
15	15:54	00700	drilling stopped 20 min. ago

BARING ("ROSMAN") SPRING (25N/34E-8N1s) WATER LEVELS

JULY 16	10:03	00706	no activity
16	13:58	00708	5 min. after purging core hole
16	14:52	00708	just finished tripping out of core hole
17	9:12	00720	stock pump on intermitt.
17	13:00	00723	"
18	12:06	00737	no activity
19	9:03	00749	"
22	15:22	00796	"
23	9:31	00805	"
23	14:36	00808	"
24	9:47	00819	" An instrument reading of 00820 corresponds to a measurement of 4.68 ft. from nail on "upper board" to spring-pool surface
24	16:27	00822	no activity--storms in area
25	10:05	00830	no activity
25	12:37	00831	purging core hole
26	10:03	00839	stock pump on intermitt.
26	14:17	00841	"
29	14:15	00860	no activity
30	9:57	00863	stock pump on intermitt.
30	13:21	00864	"
31	12:03	00866	
31	15:08	00867	stock pump on intermitt.--float of instrument may be nearing bottom of stilling well.

