

# 3D geologic map of the Deer Park 7.5-minute quadrangle, Spokane County, Washington

3D PDF INSTRUCTIONS

OBJECT DATA

- Layer001
- Layer002

No Separation

5%

50%

100%

5

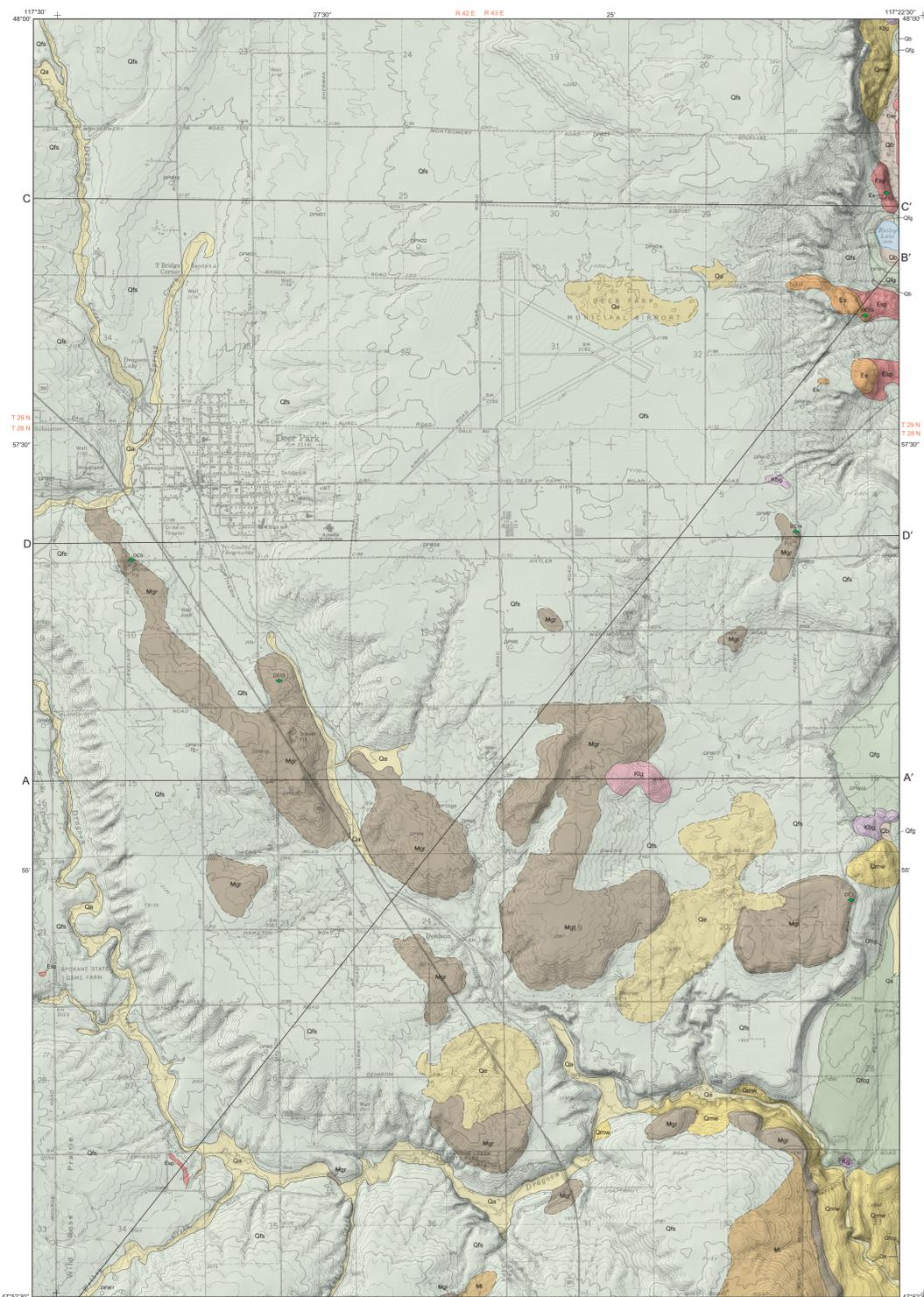
Probe

1x Z-Scale 10x Z-Scale Default Scale 3D

# Geologic Map of the Deer Park 7.5-minute Quadrangle, Spokane County, Washington

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## INTRODUCTION

The earliest geologic mapping in the Spokane area (including the Deer Park quadrangle) was by Pardee and Bryan (1926). Griggs (1966) completed a 1:125,000-scale geologic map of the western half of the Spokane 1-by-2-degree quadrangle. He later extended his mapping eastward to encompass the entire Spokane 1-by-2-degree quadrangle (Griggs, 1973). Joseph (1990) compiled a 1:100,000-scale map of the Spokane quadrangle that incorporated more detailed interpretations of Pleistocene glacial features, based on Kiver and others (1975), and basin stratigraphy, based on Swanson and others (1970). Two reports prepared for the Spokane County and the Spokane County Conservation District reported results of investigations into groundwater availability problems in Deer Park (Emcon Northwest, Inc., 1992) and water flow on Dragon Creek (Lundgren, 1998). Two cross sections from the Emcon report were modified for cross sections C and D. The right adjacent Chatwary 7.5-minute quadrangle (Derkey and others, 2005) was mapped in conjunction with this quadrangle.

## DESCRIPTION OF MAP UNITS

### Quaternary Sedimentary Deposits

- Qa Alluvium (Holocene)**—Predominantly silt, sand, and gravel deposits in present-day stream channels and on flood plains; consists of reworked glacial flood deposits (units Qfg, Qlg, and Qof) and loess; may include small alluvial fans and some mass-wasting deposits that extend onto the flood plain.
- Qb Bog deposits (Holocene and Pleistocene)**—Peat, silt, ash, mat (bog lime), and gyttja (freshwater mud with abundant organic matter) deposits (Miller and others, 1975); located predominantly in depressions over bedrock and fine-grained sediments where a drainage system is poorly developed.
- Qc Eolian deposits, predominantly dunes (Holocene to Pleistocene)**—Predominantly fine- to medium-grained sand and silt; grains are subangular to subrounded; mostly composed of quartz, feldspar, and lithic fragments; deposited as dunes and sheets; primary sand source is glacial flood deposits, predominantly sand (unit Qfs).
- Qdm Mass-wasting deposits (Holocene and late Pleistocene)**—Landslide debris with lesser amounts of debris-flow and rockfall deposits; consists mostly of a mixture of basalt blocks and Latah Formation (unit M) sediments. Basalt blocks range in size from several feet to hundreds of feet in diameter; alternatively, basalt may occur as irregularly distributed massive bodies. Mass wasting that occurred during glacial flooding incorporated flood materials. Most mass wasting occurred during or shortly after Pleistocene catastrophic floods, but some mass wasting continues to the present.
- Qgp Glacial drift (Pleistocene) (cross sections only)**—Unsorted and unstratified deposits of pebbles to large boulder-size clasts in a matrix of sand, silt, and clay. The only exposure in the Deer Park quadrangle is in a construction trench and is shown only on cross section B. Catastrophic flood events have modified surface exposures of glacial drift, resulting in removal of the matrix clasts and leaving a residuum of large boulders. This modified drift is mapped as flood gravel (unit Qfg) in the valley of the Little Spokane River in the northeastern corner of the quadrangle.

The following units are outliers flood deposits from glacial Lake Missoula. They are a composite of numerous flood events and do not represent deposits from any single flood event.

- Qfs Glacial flood deposits, predominantly sand (Pleistocene)**—Fine- to coarse-grained sand and granules with sparse pebbles, cobbles, and boulders; contains beds of glaciolacustrine deposits of glacial Lake Columbia as much as 3 ft thick and local beds and lenses of gravel composed mainly of granitic and metamorphic detritus from local sources and sources to the north; gray, yellowish gray, or light brown; subrounded to well rounded; poorly to moderately well sorted; medium bedded to massive; uneven distribution and thickness due to irregular underlying topography and varying degrees of erosion; mostly deposited when glacial Lake Missoula overflowed floods flowed into a high stand of glacial Lake Columbia.
- Qlg Glacial flood deposits, predominantly gravel (Pleistocene)**—Thick-bedded to massive mixture of boulders, cobbles, pebbles, granules, and sand; contains beds and lenses of sand and silt; gray, yellowish gray, or light brown; poorly to moderately sorted; both matrix and clast supported; locally composed of boulders and cobbles in a matrix of mostly pebbles and coarse sand; boulders and cobbles consist predominantly of local bedrock units and units found to the north and northeast of the quadrangle.
- Qlq Glacial flood-channel deposits, predominantly gravel (Pleistocene)**—Thick-bedded to massive mixture of boulders, cobbles, pebbles, granules, and sand; contains beds and lenses of sand and silt; gray, yellowish gray, or light brown; poorly to moderately sorted; both matrix and clast supported; locally composed of boulders and cobbles in a matrix of mostly pebbles and coarse sand; boulders and cobbles consist predominantly of local bedrock units and units found to the north and northeast of the quadrangle.
- Qm Grande Ronde Basalt, magnetotatigraphic units R<sub>1</sub> and R<sub>2</sub>, middle Miocene**—Dark gray to dark greenish gray, fine-grained basalt consisting of pale green augite and pigeonite grains and plagioclase laths and sparse phenocrysts in a matrix of black to dark brown glass and sparse minerals; locally vesicular with plagioclase laths tangential to vesicle boundaries; some vesicles contain botryoidal carbonate and red amorphous secondary minerals; thickness is quite variable due to irregular underlying topography. Identified in the map area on the basis of chemical analyses; between 15.6 and 16.5 m.y. old (Reidel and others, 1989).

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- M Latah Formation (middle Miocene)**—Lacustrine and fluvial deposits of finely laminated siltstone, claystone, and minor sandstone; light gray to yellowish gray and light tan; commonly weathers brownish yellow with stains, spots, and seams of limonite; poorly indurated; easily eroded and commonly blanketed by colluvium, talus, and residual soils; floral assemblages indicate a Miocene age (Kovilton, 1926; Griggs, 1976).
- Es Dacitic crystal tuff of the Sanpoil Volcanics (Eocene)**—Consists of 20 to 40 percent crystals of mostly white plagioclase (3-5 mm long) and lesser biotite (<2 mm long) in a gray, greenish gray, or brown matrix; matrix consists of devitrified glass and ash that has recrystallized to a microcrystalline mass of potassium feldspar and quartz; thickness probably does not exceed 60 ft; exposures are probable remnants of a more extensive bed contact between this unit and underlying Silver Point Quartz Monzonite (unit Esp) not observed. Whole rock analysis of samples DC58 and DC66 indicate dacitic composition.
- Esp Silver Point Quartz Monzonite (Eocene)**—Quartz monzonite consisting of distinct microporphic orthoclase phenocrysts up to 1 in. long, accompanied by smaller, zoned plagioclase, hornblende, biotite, and quartz crystals in a fine-grained groundmass; generally light gray with a greenish tinge near contact with host rocks; hornblende has a long dimension of as much as 0.4 in. and is associated with biotite; orthoclase phenocrysts are euhedral; other phenocrysts range from euhedral to anhedral, most are subhedral; titanite is the most common accessory mineral, followed by magnetite, apatite, zircon, and rare allanite (Miller and Clark, 1975); as much as 50 percent of the rock is groundmass; consists of dikes and irregularly shaped intrusive bodies. Two samples from the Chewelah 1:100,000-scale quadrangle to the north gave whole-rock Rb-Sr ages of 39.4 Ma and 46.2 Ma (Armstrong and others, 1987); recalculated K-Ar ages on rocks from the Chewelah 1:100,000-scale quadrangle were 51 Ma on biotite and 42 Ma on hornblende (Miller and Clark, 1975).
- EKgu Unknown granitic rocks (Eocene or Cretaceous) (cross sections only)**—Extensive areas of unexposed rocks shown in cross sections are believed to be granitic in composition and could be one of four known granitic rocks in the area: Eocene Silver Point Quartz Monzonite (unit Esp), leucocratic granitic rocks (unit Klg), biotite granite (unit Kbg), or Mount Spokane granite (unit Kgm).
- Kbg Biotite granite of Bear Lake (Cretaceous)**—Massive, medium- to coarse-grained, biotite-muscovite granite to quartz monzonite; consists of microcline and albite in microporphic combination; quartz, biotite, and muscovite; microcline and albite content are nearly equal; biotite and muscovite are generally less than 5 percent, muscovite subordinate to biotite. This unit may be the biotite-bearing equivalent of leucocratic intrusive rocks (unit Klg).
- Kgu Leucocratic intrusive rocks (Cretaceous)**—Medium-grained muscovite quartz monzonite; consists of microcline and albite in microporphic combination; quartz, and muscovite; microcline and albite content are nearly equal; muscovite can range up to 10 percent but is generally less than 5 percent; rarely contains a trace to 2 percent biotite; pink to cream colored; leucocratic dikes cut biotite granite (unit Kbg). Miller and Clark (1975) reported that exposures of leucocratic granitic rocks noted by Griggs (in Miller and Clark, 1975) south of Clayton (>3 mi west of the map area) were the same unit as their leucocratic muscovite quartz monzonite and that because the plagioclase is albite the rock could be classified chemically as granite.
- Kgm Biotite-muscovite granite of Mount Spokane (Cretaceous)**—Medium- to coarse-grained, massive, muscovite-biotite granite to quartz monzonite; contains medium gray anhedral quartz that commonly forms graphic intergrowths with feldspar, potassium feldspar and plagioclase are present in a ratio of about 2:3; large crystals of potassium feldspar in some exposures enclose small biotite grains; plagioclase is commonly altered; subhedral biotite comprises as much as 10 percent of the rock and forms clots; muscovite ranges from 0 to 10 percent of the rock and is present as single subhedral crystals, in clots, or with biotite; undisturbed outcrops are medium gray due to lichen cover; light gray in roadcuts and fresh exposures; weathers yellow with limonite staining. Yielded discordant K-Ar ages of 48 Ma on biotite and 53 Ma on muscovite (Miller and Engels, 1975), which are probably reset.

## GEOLOGIC SYMBOLS

- Contact—Long dash where approximately located; short dash where inferred or indefinite
- Well—Numbers correspond to well numbers on cross sections
- ◆ Whole rock geochemistry sample location—Numbers correspond to sample numbers in Table 1

## ACKNOWLEDGMENTS

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Lambert conformal conic projection  
 North American Datum of 1973, to place on North American Datum of 1983, move the projection lines 12 meters north and 79 meters east as shown by crosshair corner ticks  
 Base map from scanned and rectified U.S. Geological Survey 7.5-minute Deer Park quadrangle, 1972  
 Shaded relief generated from U.S. Geological Survey 10 meter digital elevation model; vertical exaggeration 4x  
 Digital cartography by Charles G. Carothers and Anne C. Heinze  
 Editing and production by Jareta M. Roloff and J. Eric Schuster

